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Atmospheric Secondary Organic Aerosol Formation: Parameterization of Aerosol Mass Fractions (Yields) with Confidence Intervals from Smog Chamber Data

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13 Abstract

14 A framework is presented and evaluated for the parameterization of smog chamber 15 results for use in atmospheric Chemical Transport Models (CTMs). The parameterization 16 uses an absorptive partitioning model to describe formation of secondary organic aerosol 17 (SOA). The key points of the framework are (1) the ability to fit results from several 18 types of chamber experiments; (2) the use of a basis set of surrogate compounds 19 characterized by fixed effective saturation concentrations instead of the more commonly 20 used variable saturation concentrations; (3) calculation of uncertainties of the estimated 21 SOA aerosol mass fractions outside of the fitted experimental range; (4) determination of 22 the effective enthalpy. The features of this data analysis and fitting framework are 23 demonstrated using simulated data, and actual measurements from α -pinene ozonolysis 24 experiments. Representation of SOA formation using as many as eight surrogate 25 compounds with fixed effective saturation concentrations is shown to be feasible and has 26 advantages over simpler parameterizations.

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28 Keywords: Secondary organic aerosol; modeling; parameter estimation; α-pinene

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32 1 Introduction

33 Driven by the need to simulate secondary organic aerosol (SOA) formation for 34 Chemical Transport Models (CTMs), the results of "smog" chamber studies of SOA 35 formation have been extensively parameterized. Early smog chamber experiments were 36 translated into aerosol mass fractions (yields) for models by assuming a constant aerosol 37 mass fraction for each precursor (Hatakeyama et al., 1989) or assuming independent 38 condensation of each SOA component when its concentration exceeded its saturation 39 value (Pandis et al., 1992). With additional chamber experiments and theoretical 40 investigation, it was demonstrated that SOA aerosol mass fractions of the various organic 41 precursors are not constant, but depended on aerosol concentration in a manner consistent 42 with equilibrium gas-particle partitioning (Odum et al., 1996; Pankow, 1994a; Pankow, 43 1994b).

With the completion of more chamber studies and the wider acceptance of absorptive partitioning theory for organic aerosol formation, a database of empirical concentration-dependent aerosol mass fractions was developed, and implemented in CTMs (Griffin et al., 1999; Lurmann et al., 1997; Odum et al., 1997; Strader et al., 1999). SOA studies have been reviewed by (Kanakidou et al., 2005; Seinfeld and Pankow, 2003).

50 Prior to about 2004, the main type of chamber SOA formation experiment gave 51 one aerosol mass fraction (aerosol formed / hydrocarbon reacted) after several hours of 52 oxidation and aerosol formation. These "final aerosol mass fraction" experiments were 53 limited in time resolution, usually by the time required to measure the precursor 54 concentration by GC (aerosol concentration could be monitored at minute time scales

using Scanning Mobility Particle Sizers). While real-time monitoring of VOC levels 55 56 during SOA formation was available using long path FTIR, this required high concentrations of reactants. Now that Proton Transfer Reaction Mass Spectrometry 57 58 (PTRMS) is in use in several chambers, dynamic concentrations can be measured 59 generating dozens of data points for each chamber experiment (Lee et al., 2006; Ng et al., 60 2006; Presto et al., 2005). The applicability of dynamic aerosol mass fractions (AMF) can 61 be evaluated by comparison with fixed AMF (Pathak et al., 2007). Other experiments or 62 series of experiments have been conducted to establish the relationship between 63 temperature and SOA formation. These include series of experiments with similar 64 hydrocarbon concentrations, but different temperatures (Pathak et al., 2007; Takekawa et 65 al., 2003) as well as formation at a fixed temperature followed by heating or cooling of 66 the chamber inside a temperature-controlled room or enclosure (Stanier et al., 2007). 67 This work attempts to synthesize all of these types of data to inform the SOA formation 68 parameterization.

69 Smog chamber experiments done at one temperature give no information about 70 the temperature dependence of SOA AMF. Product yields may vary as a function of 71 temperature and the vapor pressures of individual products will certainly vary with 72 In the latter case, lower temperatures should favor higher yields. temperature. 73 Temperature dependence is typically included in CTMs by calculating temperature-74 dependent saturation concentrations (or partitioning coefficients) according to the 75 Clausius-Clapeyron equation (Pankow, 1994a; Pankow, 1994b), using enthalpy of 76 evaporation ΔH for representative semivolatile species. This leads to a strong link 77 between temperature dependence of SOA concentrations and the ΔH values used.

Measurements of the temperature dependence of SOA AMF are now becoming available. One goal of this work is to present a method for including the temperature dependence results into parameter fitting, and thus reduce CTM uncertainty due to lack of knowledge of the appropriate ΔH (Pun et al., 2003; Tsigaridis and Kanakidou, 2003).

82 Two other methods for the determination of partitioning coefficients and AMF 83 The first is the direct measurement of gas and particle phase should be noted. 84 concentrations through filter/denuder GC/MS techniques (Kamens and Jaoui, 2001; Yu et 85 al., 1999). The other is a combined kinetic and thermodynamic approach where the time 86 series of aerosol mass and individual products is used to constrain a combination of rate 87 constants, stoichiometric yields, and gas-particle partitioning parameters (Kamens and 88 Jaoui, 2001). To date, these methods have been useful at elucidating gas-phase kinetic 89 mechanisms and product distributions, but have not been used to inform CTM modeling 90 As emphasized in recent overview articles on SOA of gas-aerosol partitioning. 91 formation, representation of the full complexity of SOA formation and aging will require 92 a combination of thermodynamic approaches (as emphasized in this work) as well as 93 kinetic descriptions of SOA processes (Donahue et al., 2006; Kroll et al., 2007).

Figure 1 illustrates many of the concepts to be explored in this paper. The figure 95 was created by simulating the oxidation of a VOC producing two semivolatile 96 compounds, with temperature independent stoichiometric yields. The absorptive 97 partitioning equations were solved to determine the fraction of product in the gas and 98 aerosol phases. The aerosol phase portion contributes to the aerosol mass fraction shown 99 as the z axis. The goal of this work is to determine how to parameterize the AMF surface 100 from smog chamber experimental data, including traditional final AMF, fixed-

101 temperature experiments, variable temperature experiments, and dynamic chamber 102 results. Also noted in Figure 1 is the possibility that experimental data is not available 103 across the concentration and temperature regime where the models operate. Uncertainty 104 bounds are critical for meaningful extension of the experimental results to CTM 105 applications. The uncertainty bounds need to be valid for interpolation within the 106 experimental data, and as much as possible, for extrapolation outside of the experimental 107 regime.

108 The goal of this work is to refine the estimation of parameters from chamber 109 experiments using absorption partitioning theory. In section 2.1, absorptive partitioning 110 is reviewed. In section 2.2, the parameterization algorithm is presented. Section 2.3 111 reviews the types of chamber data that are available for fitting. Section 2.4 defines 112 several terms needed for discussion of the fitting and describes four sets of pseudodata 113 used in testing the fitting algorithm. Sections 3.1-3.4 describe the application of the 114 fitting procedure to the pseudodata sets, while section 3.5 applies it to a combined α -115 pinene ozonolysis dataset including multiple chambers and multiple authors.

116 This work does not consider the parameterization of aging, heterogeneous 117 It is limited to "prompt" aerosol mass formation from reactions, or oligomerization. 118 chamber experiments, which capture chemistry and partitioning occurring over 119 approximately a 10-300 minute timescale. The choice to leave these more recently 120 discovered processes out of this work is not due to a belief that they are negligible or that 121 they do not need to be eventually incorporated into CTMs. However, the choice is 122 justifiable on at least two grounds: (1) a sound procedure for parameterization of prompt 123 SOA formation and its uncertainty is a necessary building block for more advanced

schemes; and (2) widely used CTMs (e.g. CMAQ, CAMx, and GEOS-CHEM) share the
absorptive partitioning framework, and thus can be modified without changing their basic
structure for increased accuracy and for prediction of uncertainties on SOA formation due
to limitations in the underlying chamber data.

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129 **2** Methods

130 2.1 Theoretical Basis for Semi-Empirical Fitting Equations

131 Absorptive SOA partitioning is used throughout this work, with the following 132 nomenclature. The effective saturation concentration c_i^* is defined through the 133 expression:

$$y_i c_i^* \equiv c_{i,gas}$$

135 where $c_{i,gas}$ is the gas phase mass concentration of species *i* and y_i is its mass fraction in 136 the absorbing aerosol phase ($y_i = c_{i,aer} / c_{OA}$). The variables $c_{i,aer}$ and c_{OA} are the aerosol 137 mass concentration of species *i*, and the total absorbing organic aerosol concentration, 138 respectively. This is equivalent to relating effective saturation concentration to pure 139 component vapor pressure using:

140
$$c_i^* = \gamma_i c_i^o \frac{M_{OA}}{M_i} = \frac{\gamma_i p_{L,i}^o}{RT} M_{OA}$$

141 where M_i is the molecular weight of species *i*, M_{OA} is the average molecular weight of the 142 absorbing aerosol phase, c_i^{o} is the saturation concentration and $p_{L,i}^{o}$ is the subcooled 143 liquid vapor pressure of *i*. We note two advantages for this treatment of partitioning for 144 CTM parameterizations: first, the activity coefficient is included in the effective 145 saturation concentration, and therefore the mixture is treated as pseudoideal; and second, the equations do not require explicit terms for molecular weight ratios of the precursor and products, and of the various products and the average organic aerosol molecular weight.

149 The mass fraction of species *i* partitioned to the aerosol phase, ξ_i , for a single 150 semivolatile compound in absorption partitioning is:

151
$$\xi_{i} \equiv \frac{c_{aer,i}}{c_{tot,i}} = \frac{1}{1 + \frac{c_{i}^{*}}{c_{OA}}}$$
3

where $c_{tot,i}$ is the total concentration (gas + aerosol) of species *i*. Therefore, for a mixture of semivolatile compounds, the total organic aerosol concentration is (Donahue et al., 2006):

155
$$c_{OA} = \sum c_{tot,i} \xi_i = \sum \frac{c_{tot,i}}{1 + \frac{c_i^*}{c_{OA}}}$$

156 In equations 1-4, we have not made any distinction as to the source of organic material (either $c_{tot,i}$ or c_{OA}) – it can be either preexisting in the atmosphere, or it can be generated 157 158 from reactions or from combustion emissions. Smog chambers are a special case, where 159 typically all the organic mass (for all species) is generated from oxidation of the parent 160 reactive organic gas (ROG). Some chamber experiments are conducted with preexisting 161 semivolatiles in the gas or aerosol phase, but they are the minority and would require 162 minor changes to the fitting equations and parameterization algorithm. In cases where 163 $c_{tot,i}$ is solely the result of a reaction of a precursor gas:

164
$$c_{tot,i} = \alpha_{i,molar} \Delta ROG M_i / M_{ROG} = \alpha_{i,mass} \Delta ROG$$
 5

165 In this work, all α_i values are mass-based yields. Expressing equations 3-5 for a 166 smog chamber experiment with no preexisting organic aerosol (such that $\Delta c_{OA} = c_{OA}$):

167
$$\xi \equiv \frac{\Delta c_{OA}}{\Delta ROG} = \sum \frac{\alpha_i}{1 + \frac{c_i^*}{c_{OA}}}$$
6

168 where ξ is the overall aerosol mass fraction. Equation 6 suggests that the aerosol mass 169 fraction can exceed unity, owing to increase in molecular weight upon reaction. The 170 partitioning equations (1-4) are applicable to all atmospheric absorptive partitioning 171 calculations, while the smog chamber equations (5-6) should be reserved for cases where 172 the assumption of zero preexisting adsorbing aerosol is met.

173 The term aerosol mass fraction (ζ) is used throughout instead of the term yield 174 (yield $\equiv \Delta c_{OA} / \Delta ROG$) common in SOA literature. This is done to reserve the term yield for product yields (α_i) and avoid confusion between product yields α_i and aerosol mass 175 176 fraction (AMF or ξ). For comparison to previous works, the substitution yield \equiv $\Delta c_{OA}/\Delta ROG = \xi$ can usually be made. As pointed out by Presto and Donahue (2006) 177 178 there is uncertainty on the density of organic aerosols. Aerosol mass concentration is 179 often measured by combining the measured aerosol number distribution by a scanning 180 mobility particle sizer with the assumption of sphericity and an assumed density (often 181 unity). Thus, application of equation 6 to both atmospheric prediction and to 182 experimental data reduction requires careful treatment of density. The relevant density 183 correction equations are in the appendix. In this work, all yields are calculated using assumption of aerosol density of 1 g cm $^{-3}$. 184

185 The partition coefficient $K_{om,i}$ (units $\mu g^{-1}m^3$) can be used interchangeably with the 186 effective saturation concentration, with the equation for interconversion as:

187
$$K_{om,i} \equiv \frac{c_{aer,i}}{c_{gas,i}c_{OA}} = \frac{1}{c_i^*}$$
 7

A typical procedure (Cocker et al., 2001; Griffin et al., 1999; Odum et al., 1996) for data reduction of chamber experiments is to use equation 6 to regress smog chamber data (pairs of c_{OA} and ζ) to fit either 2 parameters (e.g., a single α_i and saturation concentration c_i^*) or 4 parameters (e.g., a pair of α_i values with corresponding saturation concentrations).

When aerosol mass fractions are needed at different temperatures, the Clausius-Clapeyron equation is used to adjust the effective saturation concentrations.

195
$$c_i^*(T) = c_i^*(T_{ref}) \frac{T_{ref}}{T} \exp\left[\frac{\Delta H_i}{R} \left(\frac{1}{T_{ref}} - \frac{1}{T}\right)\right]$$
8

196 where T_{ref} is a reference temperature where a reference effective saturation is defined, and 197 ΔH_i is the enthalpy of evaporation. There is considerable uncertainty regarding the value 198 of ΔH to use in this equation (Bian and Bowman, 2002; Stanier et al., 2007; Strader et al., 199 1999; Tsigaridis and Kanakidou, 2003). One approach (which usually gives to strong of 200 a temperature response) is to use the few known ΔH values of identified SOA 201 components.

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203 2.2 Parameterization Algorithm

The parameterization strategy is to determine values of c_i^* , α_i , and ΔH_i so that aerosol mass fractions (AMF) predicted by equations 6 and 8 match experimental values as closely as possible. This is done by nonlinear least squares regression. The actual regression is straightforward, but there are several challenging issues related to the parameterization: (1) selecting the number of basic set compounds (*n*) to use, (2) deciding whether to use fixed or variable effective saturation values, (3) estimating the temperature sensitivity accurately (or at all), and (4) estimating confidence intervals across the range of c_{OA} and temperature values anticipated in the CTM application. An overview of the procedure is shown in Figure 2.

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214 <u>2.2.1 Selection of extreme values of c^* </u>

Equations for selecting extreme values of c^* are derived in the appendix. They depend on the accuracy goals, organic aerosol concentrations, and temperatures of the expected CTM application. For c^* values referenced to 298 K (25°C) the equations are:

218
$$c_{\min,298K}^* \approx \delta \exp\left[\frac{\Delta H}{R}\left(\frac{1}{T_{\max}} - \frac{1}{298}\right)\right] \qquad 9$$

219
$$c_{\max,298K}^* \approx \frac{5}{f} c_{OA,\max} \exp\left[\frac{\Delta H}{R} \left(\frac{1}{T_{\min}} - \frac{1}{298}\right)\right]$$
10

where δ is the allowable error (µg m⁻³) in the semivolatile partitioning calculation under clean conditions, T_{min} and T_{max} are the temperature limits of CTM modeling application, and *f* is the maximum fractional error in the parameterization at high concentration (e.g. 0.1 for ±10% allowable organic aerosol concentration). Some values from equations 9 and 10 are given in Table 1 for $\Delta H / R$ of 12,000 K. For the α -pinene cases examined in the results and discussion section, it turns out that this choice is overly conservative, and a value of 4,000 K will suffice.

227

228 <u>2.2.2 Selection of number of *c** values</u>

A priori specification of the number of c^* values (including c^*_{min} and c^*_{max}) required for a suitable fit and for confidence interval characterization is difficult. A practical approach is to space the c^* values on a lognormal basis (e.g. 0.01, 0.1, 1, ... 10^3 , 10^4 , $10^5 \ \mu g \ m^{-3}$) and thus the number of elements in the basis vector is determined by the spacing between c_{min}^* and c_{max}^* . The regression calculation is quick and can be repeated easily with a different basis, so trial and error selection of the c^* basis set is feasible. From doing the fits for this work, we recommend:

236
$$n \approx a \log_{10} \left(c_{\max}^* / c_{\min}^* \right) + 1$$
 11

where *a* is from 0.5 to 1.0. There exist cases where the *experimental data* can be fit with a small basis set (e.g. n = 1 or 2). This will be true especially when the data have a small dynamic range in terms of c_{OA} values. When equation 11 gives a large recommended basis set but one or two well selected c* values are sufficient to fit the experimental data, the additional parameters are still necessary for calculating uncertainty when the parameterization is extrapolated outside the range of the underlying experimental data.

Although the current work is limited to fitting AMF (and not a full mass balance for the SOA formation chemical reaction) the use of a suitably large basis set may have further utility in tracking the mass balance, secondary reactions, aging, and atmospheric fate of biogenic and anthropogenic ROG emissions (Donahue et al., 2006; Robinson et al., 2007).

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249 <u>2.2.3 Regression</u>

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The objective function minimized during regression is

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$$OF = \sum_{i=1..m} \left[\xi - \hat{\xi} \left(\alpha_i, \Delta H, c_i^*(T_{ref}), T_{ref}, c_{OA}, T \right) \right]^2$$
 12

where ξ are the measured AMF (*m* data points) and $\hat{\xi}$ are the modeled AMF calculated using equations 6 and 8. The modeled AMF are calculated using fixed parameters (T_{ref}), and experimental variables c_{OA} and T. The adjustable parameters are α_i , a single ΔH (instead of $n \Delta H_i$ values). Fixed saturation concentrations c_i^* are used in this work, but they can also be adjustable parameters if n is small.

257 The fact that the AMF is linear in the adjustable parameters α_i (Equation 6) 258 greatly speeds up the calculations. At any step in the regression (and for any value of 259 ΔH), the best set of α_i can be computed by solving the following equation subject to yield 260 non-negativity constraints:

261
$$\begin{bmatrix} \frac{c_{OA,1}}{c_{OA,1} + c_1^*} & \frac{c_{OA,1}}{c_{OA,1} + c_2^*} & \cdots & \frac{c_{OA,1}}{c_{OA,1} + c_n^*} \\ \vdots & & \vdots \\ \frac{c_{OA,m}}{c_{OA,m} + c_1^*} & \frac{c_{OA,m}}{c_{OA,m} + c_2^*} & \cdots & \frac{c_{OA,m}}{c_{OA,m} + c_n^*} \end{bmatrix} \begin{bmatrix} \alpha_1 \\ \alpha_2 \\ \vdots \\ \vdots \\ \alpha_n \end{bmatrix} = \begin{bmatrix} \xi_1 \\ \xi_2 \\ \vdots \\ \xi_m \end{bmatrix}$$
262
$$\alpha_i \ge 0$$

Because of this feature, nonlinear minimization is performed for only ΔH . In this work, the Levenberg-Marquardt method implemented in MATLAB's *lsqnonlin* function (MathWorks, 2002) is used for the nonlinear minimization and the MATLAB function *lsqnonneg* is used for the nonnegative least squares problem. The regression calculation is not sensitive to initial guesses and takes a few seconds for fixed c_i^* values, with n=8and hundreds of datapoints. The goodness of fit will be characterized using the metrics shown in Table 2.

Initial confidence intervals are calculated using the asymptotic confidence interval method (Seber and Wild, 2003) as implemented in MATLAB's *nlpredci* function. The uncertainty bounds are calculated *without* using the constraints that there are physical 273 upper and lower limits to AMF. Therefore, far enough from the data, the inherent 274 multicollinearity in the basis set manifests itself as unreasonably high and low values of 275 the confidence intervals. These are truncated using the logic that AMF cannot decrease 276 with increasing c_{OA} (at constant temperature), that AMF cannot increase with decreasing 277 c_{OA} (at constant temperature), and that AMFs have both lower (zero) and upper (~1.5) 278 limits.

279 This approach often leads to wide nonphysical confidence intervals when 280 extrapolating outside the experimental range. Therefore, a second method of calculating 281 confidence intervals is employed: repeated random (Monte Carlo) sampling of selected α_i 282 values. The basic idea is that the experimental data may poorly constrain certain yields 283 associated with low volatility (low c*) and high volatility (high c*) elements in the basis 284 Therefore, the yields for these compounds are selected randomly while the set. 285 remaining α_i 's are fit to the experimental data. The algorithm for the Monte-Carlo 286 confidence intervals is included in an appendix. Random sampling is done in large 287 (~1000) simultaneous trials with either 1, 2 or 3 random α_i 's. Random sampling is 288 stopped when change in the confidence intervals is smaller than a preset tolerance. This is 289 the slowest part of the overall parameterization code. 1000 trials require \sim 1 second using 290 MATLAB's *lsqnonneg* routine on 2.1 MHz PC with 2 GB RAM (this is while estimating 291 confidence intervals with 300 points in the c_{OA} vs. T space). The trials need to be 292 repeated from 2-20 times for each value of ΔH explored, so exploration of the confidence 293 intervals using 10 values of ΔH can take 1-2 CPU minutes. The combination of these 294 techniques for different cases will be discussed in subsequent sections.

296 2.3 Description of Data Used in Fitting

297 Data used in this work to parameterize SOA AMF fall into three categories. The 298 first category is the traditional fixed-temperature smog chamber AMF experiment. In 299 these experiments, the AMF of secondary organic aerosol (defined as the mass of created 300 aerosol divided by the mass of reactant consumed) is measured after the reactive organic 301 The relative humidity and concentration of gas is oxidized in a large chamber. 302 preexisting aerosol is controlled. Temperature may be controlled or uncontrolled, 303 although a single temperature is usually recorded corresponding to the temperature at the 304 time when the AMF estimate is made. A series of experiments is usually completed 305 according to a design to provide variation in reactant concentration, c_{0A} , and or 306 temperature. Oxidation is by OH, O₃, or NO₃. O₃ can be injected directly from a corona 307 discharge source, or photochemical oxidation can be initiated from NO_2 photolysis. If 308 ozonolysis is the targeted process, an OH scavenger such as 2-butanol can be included in 309 the chamber in excess. As discussed in the review by Kanakidou et al. (2005) (Kanakidou 310 et al., 2005), there is a large body of literature data based on experiments of this type for 311 various precursor-oxidant combinations. These types of experiments are referred to as 312 static, final AMF, or traditional smog chamber experiments.

A second category of data available for fitting is from chamber experiments where VOC concentration can be monitored rapidly using PTRMS and aerosol concentrations are measured at minute time scale by SMPS (Presto and Donahue, 2006; Presto et al., 2005). Therefore, a time series of c_{OA} , ΔROG , AMF, and temperature is generated. 318 A third category of data available for fitting is from chamber experiments that 319 include a change in temperature following aerosol formation. This is done solely for the 320 purpose of assessing the temperature sensitivity of aerosol partitioning. Rather than 321 compare two experiments at different temperatures (each with independent uncertainties 322 that complicate attributing changes in AMF to ΔT), the ramp or step change in 323 temperature allows direct observation of the change in partitioning. The technique is 324 called TREVA (Temperature-Ramped Equilibrium Volatility Analysis). Results are 325 available only for the α -pinene-ozone and β -pinene-ozone systems (Pathak et al., in press, 326 2007; Pathak et al., 2007; Stanier et al., 2007).

327 Data from all three of these experiment types is used in the final section of the 328 results and discussion to parameterize AMF and uncertainty for α -pinene ozonolysis. 329 The specific datasets are listed and referenced in Table 3. Individual datapoints used in 330 fitting are listed in a supplemental data section.

331

332 2.4 Nomenclature and Description of Synthetic Data Used in Fitting

333 Synthetic data, also called pseudodata, is useful for demonstrating that the fitting 334 procedures work and assessing specific strengths and weaknesses of the parameterization 335 strategy. Values of α_i , c^*_i , and ΔH_i are used to simulate AMF experiments, generating 336 synthetic (pseudo) data as datapoints with values of c_{OA} , ξ , and T. These can be 337 simulated with random error and bias if desired. We will use the following terminology: 338 • *Underlying model*: the α_i , c^*_i , and ΔH_i used to generate synthetic data

- 550
- *Ideal AMF*: ξ calculated from the underlying model
- 340

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• Synthetic data or pseudodata: data (ξ , temperature, and c_{OA}) used in regression

341	• <i>Predictions</i> or <i>regression predictions</i> : AMF $(\hat{\xi})$ calculated according to the
342	regression output. For pseudodata cases, regression predictions are compared to
343	synthetic data, where agreement is expected, and to ideal data, where agreement is
344	not necessarily expected and will depend on many factors.
345	• <i>Regression model:</i> the structure (equation 6) and parameters for
346	$\hat{\xi}(\alpha_i, \Delta H, c_i^*(T_{ref}), T_{ref}, c_{OA}, T)$. The adjustable parameters are α_i (<i>n</i> values) and ΔH .
347	T_{ref} is a fixed parameter (set to 298 K for this work). c_{OA} and T are the
348	independent experimental variables.
349	• <i>Basis set</i> or <i>basis vector</i> : the fixed values of c _i * used in a regression model.
350	• Parameterization applicability domain (also model domain): the expected range
351	of c_{OA} and temperatures where the parameterization will be employed.
352	• <i>Experimental domain:</i> range of c_{OA} and temperatures in the experimental or
353	synthetic data.
354	• <i>Temperature sensitivity:</i> the change in aerosol concentration with temperature due
355	to semivolatile partitioning, expressed as $\partial \ln(c_{OA})/\partial T$ with units of K ⁻¹ . For SOA
356	from α -pinene ozonolysis, this is expected to be ~ -0.004 to -0.036 K ⁻¹ (Stanier et
357	al., 2007). For consistency, all temperature sensitivity values in this work are
358	calculated at 40 μ g m ⁻³ and 20°C.
359	The ideal models and synthetic data experiments used in this work are described in Table
360	4. Two underlying models are used to generate four sets of pseudodata.
361	A 2-product underlying model is used to generate sets A, B1, and B2. In
362	pseudodata set A, the sequence of hypothetical experiments includes 19 chamber final
363	AMF experiments, with experimental c_{OA} values from 7-525 $\mu g\ m^{-3}$ and temperatures

from 26 to 37°C. Experiments are assumed to have very low independent random error in determination of Δ ROG such that the coefficient of variation (CV= σ /mean) on Δ ROG from repeated smog chamber experiments would be 0.007.

367 In pseudodata set B1, 21 final AMF experiments are simulated and they are 368 assumed to come from 3 separate series of chamber experiments. The first series of 369 experiments is 10 experiments, all at 26°C. They are assumed to have random error such 370 that CV=0.02, but they are also assumed to have a series-specific bias of +10% in 371 measured AMF (e.g. true AMF of 0.2 would be sampled as 0.22). The second series of 372 experiments is 6 experiments, all at 37°C. They are assumed to have random error such 373 that CV=0.02, but they are also assumed to have a series-specific bias of -10% in 374 measured AMF. A third series of experiments is 5 experiments at the same fixed ΔROG , 375 but each at a different temperature, from 17-37°C. They are assumed to be free of 376 systematic error but have random error such that CV=0.02.

377 In pseudodata set B2, the pseudodata experiments from B1 are supplemented with 378 5 temperature ramp experiments, where the AMF is measured at 5 different temperatures 379 as the chamber is heated or cooled. The data is assumed to be noisy, and to have 380 correlated errors. Specifically, each of the 5 experiments is assumed to have a bias 381 shared by all 5 datapoints, associated with the determination of the ΔROG for that 382 experiment. This is generated randomly such that the CV=0.06. The details of the 383 hypothetical experiments are that each experiment gives 5 AMF datapoints at 384 temperatures ranging from 22 to 39°C. Each datapoint was generated with an additional 385 independent error, corresponding to a CV=0.03.

386 In pseudodata set C, a much more complex physical mixture is assumed, with 14 products with c* values ranging from $\sim 10^{-2} \ \mu g \ m^{-3}$ to $\sim 10^4 \ \mu g \ m^{-3}$. No single SV 387 component was dominant, but the most prominent three had yields (α_i) of 0.09, 0.042, 388 and 0.03 at c* values of $10^{-0.2}$, $10^{0.7}$ and $10^{2.6} \ \mu g \ m^{-3}$, respectively. Two of the fourteen 389 390 had yields of less than 0.005. A full list is given in table 4. The pseudodata comprising 391 dataset C are sampled in a very similar sequence of experiments as in set A. 19 simulated 392 experiments, with low random error (CV=0.007), no systematic error, spanning min/max c_{OA} values of 7 and 525 µg m⁻³, and including temperatures from 26 to 37°C. 393

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3 Results and Discussion

396 3.1 Ability to Fit Synthetic Data

Following the flowchart in figure 2, there are 3 separate steps before any calculations begin. The first requirement is experimental data. For this demonstration, pseudodata set A is used, where 19 pseudo-experimental data points have been sampled from an underlying 2 product model (see also section 2.4).

401 The second requirement is a basis vector of c* values to be used in the fit; the 402 basis vector needs to be consistent with the CTM model domain, as described in section 2.2. For this demonstration, we assume the model domain is: 0-40°C; $c_{OA,max}$ of 25 µg m⁻ 403 ³, acceptable error (δ) of 0.07 µg m⁻³ in SOA partitioning under very clean conditions, 404 and an acceptable error of 1.25 µg m⁻³ ($f c_{OA max}$ with f=0.05) in SOA partitioning under 405 406 very polluted conditions. This set of values (listed in Table 1) is requires (equation 9 and 10) c_{min}^* of 0.01 µg m⁻³ and c_{max}^* of 10⁵ µg m⁻³. Equation 11 puts *n* at 4-8. The choice of 407 8 corresponds to a decadal spacing $(0.01, 0.1, 1, 10, 100, 10^3, 10^4, 10^5 \,\mu\text{g m}^{-3})$. 408

409 The third requirement is to form vectors to represent the model domain as a $p \ge i$ 410 grid of c_{OA} and temperature values. The grid needs to have enough points so that 411 curvature in the AMF surface can be suitably represented. In this work, we use 20 values 412 of c_{0A} (p=20) and repeated all calculations at 5 different temperatures (0, 10, 20, 30 and 413 40°C). Therefore, an output of the algorithm will be AMF estimates and confidence 414 intervals at the p x j points. Finally, a vector of ΔH values is required for probing the uncertainty limits. In this work, we test values ranging from 10 to 120 kJ mol⁻¹ in 5-15 kJ 415 mol⁻¹ increments. For most datasets, several of the extreme values cannot fit the data 416 417 under any choice of α_i . Therefore, for a final fit, a narrower range of more tightly spaced 418 ΔH vector may give the widest uncertainty bounds.

419 With those three items set, the optimization and the determination of the confidence intervals can proceed. The result for the basis c* vector (0.01, 0.1, 1, 10, 100, 420 10^3 , 10^4 , and $10^5 \ \mu g \ m^{-3}$) is shown in Figure 3. The regression lines (thick solid lines) go 421 422 through the data very well (the mean absolute fractional error is less than 1%). The 423 uncertainty bounds (grey) are narrowest where there is the most data (30°C) and become 424 wider away from the data. The underlying model included two SOA species with ΔH values of 80 and 60 kJ/mol. The regression estimated a ΔH value of 77 kJ mol⁻¹. The 425 426 fitted α values were (0; 0; 0.027; 0.097; 0; 0.069; 0.25; 0.25). The ideal AMF (narrow 427 dotted lines) are encompassed within the confidence intervals for nearly the entire 428 parameterization applicability domain (88 percent of the ideal AMF fall within the 429 confidence intervals). The uncertainty bounds are output not as a parameterized function, 430 but rather as an array with a value for the upper and lower CI at each of the p x i aerosol-431 temperature combinations. The confidence intervals are small in parts of the model 432 domain, reflecting the high precision in the pseudodata. In Figure 3, and in all 433 subsequent figures of the same style, the experimental data is shifted to the nearest 434 temperature line using a preset temperature sensitivity of -0.015 K⁻¹. For fitting, the data 435 is fit at its actual temperature – the shift is only so that the data can be summarized in one 436 AMF vs. c_{OA} plot.

437

438 3.2 Exploring Goodness of Fit vs. Basis Set Selection

To explore the relationship between the choice of the basis set and the quality of the parameterization, we now repeat the fit of pseudodata set A with different basis vectors, varying the number of elements in the basis vector, their spacing, and the extreme values of the basis set (c_{min}^* and c_{max}^*). The pseudodata are unchanged; only the basis vectors are changed. The choices are then compared on the basis of goodness-offit, fraction of data within confidence intervals, and confidence interval width.

445 The results are shown in Figures 4 and 5. Five different fits with basis vector 446 length 2 (n=2 fits) were done. These are denoted in Figure 4 as entries 2a-2e. Four 447 different n=3 fits were included in the comparison (3a-3d), as well as 2 n=4 fits (4a, 4b). 448 Single fits were done at n=5, 6, 7, 8, and 9, respectively. The n=5-7 cases had lognormal spacing between 0.01 and 1,000 μ g m⁻³. The n=8 fit is repeated from section 3.1. The 449 n=9 fit has lognormal spacing from 0.01 to c_{max}^* to $10^5 \,\mu g \,m^{-3}$. Figure 4 has performance 450 statistics for all 16 of these cases. Figure 5 shows AMF vs. c_{OA} plots for 4 of the 16 451 452 cases.

453 On the y axis of figure 4a is a goodness-of-fit of the regression prediction relative 454 to the synthetic data (black bars) and to the ideal AMF (white bars). The plotted error is

455 the mean error fraction, err_{frac} (see table 3). For example, a prediction of 0.18 AMF 456 versus a data point of 0.20 is expressed as a 10% error in Figure 4a. Figure 4b graphs the 457 percentage of the true AMF across the entire domain that fall within the confidence 458 intervals. Figure 4c shows the width of the confidence intervals at 4 specific points in the system. One point (30°C, 26 μ g m⁻³, triangle symbols) is right in the heart of the 459 460 pseudodata. One would expect the narrowest CI for that portion of the regression. Another, $(30^{\circ}C, 7 \text{ µg m}^{-3}, \text{ circle symbols})$ is within the data, but just barely. Two points 461 represent the extrapolation in temperature (10°C, 26 µg m⁻³) and in temperature and 462 concentration (10°C, 7 μ g m⁻³). 463

464 Several conclusions can be drawn from Figures 4 and 5. First, low n (2-3) 465 components in the basis vector) fits are variable in goodness-of-fit relative to the data. 466 As shown in Figure 4a (fits 2a-3d), some basis sets lead to good fits, while other fail 467 badly. On the other hand, using an n=2 basis vector for estimating confidence intervals on the regression gives poor results regardless of the c* values. The fraction of ideal 468 469 model points falling in the confidence intervals (Figure 4b, fits 2a-2e) is fairly low, while 470 the size of the confidence intervals is highly dependent on the goodness of fit, varying 471 from unreasonably narrow (2a) to very wide (2c). The variability in the regression results 472 is reinforced by the n=2 and n=3 panels of Figure 5; the goodness-of-fit and the width of 473 the confidence intervals show extreme variability, depending on the exact choice.

In case 2a, the c* values selected correspond exactly to the values in the underlying model. Therefore, the data are fit very well, giving the lowest value of the objective function of all the cases in figures 4 and 5. However, the excellent fit coupled

with the fact that there are no poorly constrained regression parameters gives very narrowconfidence intervals.

479 The second point to draw from figures 4 and 5 is that the regression results 480 improve in quality and consistency when a basis vector with $n \ge 6$ and appropriate c_{min}^* and c_{max}^* are used. The exact choice of individual c^* values becomes much less 481 482 important as n increases to values of 6 and higher. In all the $n \ge 6$ cases, the errors 483 relative to the data and pseudodata are small. The fraction of ideal points within the 484 confidence intervals are reasonable, and the uncertainty in the prediction is smallest in the 485 heart of the data, and grows larger as the degree of extrapolation increases. AMF vs. c_{OA} 486 plots can be seen with uncertainty bounds for the n=6 case (Figure 5) and the n=8 case 487 (Figure 3).

488

489 3.3 Temperature Sensitivity and ΔH

490 In some fitting applications the data will constrain the temperature sensitivity of 491 aerosol concentrations. In other words, in a plot such as Figure 3, the AMF will be 492 equally well predicted at a range of different temperatures. Pseudodata set A is, for the 493 most part, able to constrain the change in AMF with temperature. For example, the n=7 regression recovers a temperature sensitivity of -0.022 K⁻¹. The temperature sensitivity 494 of the underlying model at the same point is -0.021 K⁻¹. The reason for this good 495 496 agreement is that the pseudodata includes data from a range of temperatures, there is no 497 correlation between error and temperature, and random errors are small. The fitted ΔH value (77 kJ mol⁻¹) is in agreement with the underlying model (60 and 80 kJ mol⁻¹), as 498 499 explained in section 3.1. Furthermore, in pseudodata set A, the ΔH value is tightly

500 constrained, with a standard error of \sim 3 kJ mol⁻¹ and visible deterioration of the goodness 501 of fit when Δ H is reduced from 77 to 60.

In still other scenarios the data will *not* constrain temperature sensitivity. This can occur when all data were sampled within a narrow temperature range. In that case, the confidence intervals will be narrow at temperatures with data, and increasingly wide at other temperatures. Or the cause can be excessive random error in the data masking any temperature signal. This should lead to wide confidence intervals at all temperatures. Another possibility is that correlations between error and temperature lead to an *incorrect* apparent temperature sensitivity in the data.

509 To illustrate the case of incorrect apparent temperature sensitivity, a pseudodata 510 set B1 was generated. In pseudodata B1, a hypothetical series of experiments at 29 °C 511 has AMF 10% higher than their true values (e.g. 0.11 instead of 0.10). Another 512 hypothetical series of experiments at 37 °C reports AMF 10% lower than their true 513 values. The underlying model has SOA Δ H values at 80 and 60 kJ/mol, and the true underlying model temperature sensitivity is -0.021 K^{-1} . The regression result is shown in 514 515 Figure 6a. As expected, the regression predictions are incorrect in terms of temperature sensitivity; the recovered ΔH value is too low (40 kJ mol⁻¹) and the temperature 516 517 sensitivity of the regression prediction is -0.019 K⁻¹.

518

519 <u>3.3.1 Using "Temperature Ramp" Chamber Experiments</u>

520 One method of overcoming poorly constrained temperature sensitivity is through 521 repeated smog chamber experiments at different temperatures (Pathak et al., 2007). 522 However, the number of experiments may need to be fairly large (>15) to overcome

523 random error between chamber experiments. Another type of experiment design is a 524 temperature ramp design, where a temperature controlled smog chamber is heated or 525 cooled quickly (minutes) after initial SOA formation (Stanier et al., 2007). This has the 526 advantage of minimizing the effect of experiment-to-experiment noise in measuring the 527 temperature sensitivity. This type of experiment was included in pseudodata set B2. Set 528 B2 is identical to B1 except that 5 temperature ramp experiments, each generating 5 529 datapoints, are added to the 19 experiments already in set B1. Figure 6b shows how the 530 regression result when set B2 is fit. The temperature ramp experiments, although noisy 531 compared to the original data, meet the objective of improving the estimation of the 532 temperature sensitivity. The regression shown in figure 6b has a temperature sensitivity of -0.022 K⁻¹ and a regressed Δ H value of 67 kJ/mol. 533

534 Similar results to figure 6 were obtained by reversing the biases in the 29 and 535 37°C data series. In the above paragraphs and in Figure 6, the lower temperature data 536 was biased high while the higher temperature data was biased low, leading to an apparent 537 low temperature sensitivity. If the biases are reversed, the apparent temperature 538 sensitivity becomes large compared to the ideal temperature sensitivity of the underlying 539 model. For example, with the true temperature dependence the same as above (-0.021 K^{-1}) ¹), the initial regression's value is -0.034 K^{-1} . Once the temperature ramp data is added, 540 the sensitivity goes to -0.024 K^{-1} . 541

542 Volatility TDMA experiments (An et al., 2007; Offenberg et al., 2006; Philippin 543 et al., 2004) may also be suitable for this type of analysis, as long as equilibrium is 544 achieved in the flow setup.

546 <u>3.3.2</u> User Specified Limits on Temperature Sensitivity

If the CTM modeling results are sensitive to temperature effects on partitioning, 547 548 then realistic temperature sensitivity and confidence intervals are highly desirable. If no 549 experimental data is available as a constraint, then suggested limits are (based on α pinene SOA) $\partial \ln(c_{0A})/\partial T$ of 0 to -0.04 K⁻¹ (Stanier et al., 2007). Bracketing temperature 550 551 sensitivity is preferred over the simpler approach of just limiting the ΔH value used 552 equation 6 to a preset range. The reason for this is that the value of ΔH required for a 553 certain temperature sensitivity depends on the basis vector selection and aerosol 554 concentrations. At least two methods exist for using artificial limits on temperature 555 sensitivity. One is to simply check confidence intervals for adherence to the user selected 556 temperature sensitivity, and to narrow confidence intervals appropriately. The other is to 557 include artificial data in the fitting itself.

558

559 3.4 Pseudodata Set C – Complex Mixture

560 The underlying model for sets A and B was relatively simple, with only 2 SOA 561 products. Actual organic aerosols, even for single precursors, are much more complex. In case C, the SOA mixture is assumed to have 14 components with c* values ranging 562 from $\sim 10^{-2} \ \mu g \ m^{-3}$ to $\sim 10^4 \ \mu g \ m^{-3}$. No changes to the algorithm or procedures are 563 564 required. The result of the fitting algorithm is shown in Figure 7. A basis vector of 8 lognormally space c* values (from 0.01 to $10^5 \ \mu g \ m^{-3}$) was used to perform the 565 566 The fit results (error relative to data, relative to underlying model, regression. 567 temperature sensitivity, and confidence intervals) are quite good. The mean absolute 568 fractional error for prediction of the datapoints is less than 0.02. All of the ideal model values are within the confidence intervals. The underlying model temperature sensitivity is -0.026 K⁻¹ and the regression temperature sensitivity is -0.022 K⁻¹ (with a fitted Δ H of 571 56 kJ mol⁻¹).

572 One-to-one mapping of the ideal to the fitted yields (α values) is not expected at 573 the individual basis vector level. Nevertheless, broad agreement is seen when ideal and 574 fitted yields are compared. When mapped to the mapped to the nearest c* value, the true 575 yields would be: [0.002 0 0.04 0.07 0.04 0.11 0.02 0]. The corresponding fitted values 576 are: [0 0 0.09 0.03 0.09 0.14 0 0]. Grouping into four bins instead of eight, the 577 comparison is closer: true yields at [0.002 0.11 0.15 0.02] and fitted yields at [0 0.12 578 0.14 0].

579

580 3.5 Fitting α -pinene ozonolysis data

We now demonstrate the fitting algorithm on real data. The fitting is very similar to the pseudodata cases presented above. The main differences are that the dataset is larger, that the data somewhat noisier, and that new types of possible input data are included (final AMF experiments, dynamic AMF experiments using PTRMS, and temperature ramp experiments).

586

587 <u>3.5.1 Demonstration algorithm on large α -pinene ozonolysis dataset</u>

The data used for fitting is summarized in Table 3, and individual datapoints are listed in the appendix. For the model applicability domain the "general" model case from Table 1 was selected (0.1-60 μ g m⁻³ with a temperature range from 0-40°C). While table 1 (based on equations 9-11) recommends a c*_{min} and c*_{max} of 1x10⁻³ and 2.4x10⁵,

respectively, these are calculated based on a single component with ΔH of 100 kJ mol⁻¹ 592 593 (see appendix for derivation), and the real aerosol does not have nearly that temperature 594 sensitivity. Therefore, a narrower range might be employed. This necessary range was tested by fitting with basis vectors of length 9 (10^{-3} to 10^{5}), length 7 (10^{-2} to 10^{4}), and 595 length 5 (10^{-1} to 10^{3}). The fits and confidence intervals were nearly identical for the n=7 596 597 and n=9 cases. However, the confidence intervals change somewhat in the n=5 case. 598 Specifically, the confidence intervals narrowed at low and high concentrations from the removal of the 0.01 and 10,000 μ g m⁻³ entries in the basis set. Therefore, all of the fits in 599 section 3.5 utilize an n=7 basis set with decadal c* values from 10^{-2} to 10^4 µg m⁻³. 600 Because many of the experimental c_{OA} values are greater than the upper limit of the 601 model applicability domain, the plots for this section will be extended to 200 μ g m⁻³. 602

603 The fitting result for the n=7 case is shown in Figure 7. The mean absolute 604 fractional error is 0.08. In other words, the average error on an AMF of 0.2 is ± 0.016 . 605 From the plot, it can be seen that the source of this fitting error is the variability in the 606 data itself. There is obvious variability in the experiments, and some bias between the different series of experiments. The yield values (for $c^* = 10^{-2}$ to $10^4 \ \mu g \ m^{-3}$) are [0 0 607 0.073 0.072 0.087 0.29 0.29] where the 0.29 values at α_6 and α_7 are user-selected upper 608 609 limits (the data is not sufficient to constrain the most volatile component yields). Fitted ΔH is 33 kJ mol⁻¹ and temperature sensitivity was -0.02 K⁻¹. The largest confidence 610 intervals occur under two conditions: at less than 1 $\ \mu g \ m^{-3}$, and also at temperatures 611 612 below 20°C. The cause of both of these areas of large uncertainty is a relative scarcity of 613 data.

614 <u>3.5.2</u> Removing correlated data in highly time-resolved data

615 The use of PTRMS to track the reacting VOC concentration, and the use of SMPS 616 or aerosol mass spectrometers can generate multiple data points per experiment. The two 617 PTRMS experiments fit in this section have 38, and 89 datapoints, respectively. The 618 Durbin-Watson test, a test for serial correlation in the residuals, is performed by first 619 fitting a basis vector of length 7 to the data for the PTRMS experiment in question only. 620 The residuals from this fit are used in the Durbin-Watson test (Hamilton, 1992). In one 621 of the 2 cases, the 38 residuals were determined to not be correlated. In the other, the 622 Durbin-Watson test did indicate excessive correlation. This was removed by averaging 623 successive data points (starting with those contributing most to the Durbin-Watson score) 624 until either the test was passed, or the number of data points was reduced to 15.

625 3.5.3 Response of confidence intervals to key data

626 Figure 9 shows the result when only a subset of data is fit. The data that is used is 627 that of Cocker, 2001. This data was selected because all of the samples are at 29±1°C. Therefore, not surprisingly, extrapolation to different temperatures gives significant error. 628 The best fit Δ H and temperature sensitivity for this dataset are 10 kJ mol⁻¹ and -0.010 K⁻¹. 629 630 The uncertainty bounds are determined by the range of allowed ΔH values in the Monte Carlo sensitivity study. Those shown are for limits of 5 and 120 kJ mol⁻¹. A wider 631 632 allowed range would give even wider confidence intervals for the temperatures other than 29±1°C. For example, setting ΔH to 120 kJ mol⁻¹ gives a temperature sensitivity of in 633 this case at 20°C and 40 μ g m⁻³ of -0.046 K⁻¹. 634

Two conclusions can be drawn from figure 9. First, the confidence interval algorithms used here respond in the correct qualitative fashion as key data are added/subtracted to the dataset. Second, extending the temperature range and/or temperature sensitivity of AMF measurements is just as necessary for reduction in overall
uncertainty as is extending the concentration range. Figures 8 and 9 are directly
comparable and show the effect of the increase in the range of data on the uncertainty
bounds.

642

643 3.6 Discussion of application to Chemical Transport Modeling

644 The paper and algorithm have focused on leaving the basis set selection flexible, to be 645 varied as needed according to the demands of the expected model application (e.g. T_{min}, 646 T_{max}, c_{OA.min}, c_{OA.max}, and allowable errors in partitioning). If followed literally, that idea 647 could lead to a proliferation of data reductions, each tailored to a specific model domain; 648 that is not a direction advocated by the authors. Rather, the key concepts from this work 649 that should translate to improved modeling are (1) mismatches in the fitting basis vector 650 and model organic aerosol concentrations and temperatures can have undesirable 651 consequences; (2) realistic confidence intervals on aerosol mass fraction are possible; (3) 652 multiple experiment types can be integrated to give improved parameterization; (4) 653 comparisons of aerosol mass fractions (measurements, predictions, and confidence 654 intervals) and temperature sensitivity are much more informative that comparisons of regression parameters (α_i , c^*_i , and ΔH); and (5) high *n* fixed c^* basis vector fits are 655 656 feasible and useful.

A practical method of achieving the improved modeling results may be for wider adoption of a decadal basis set including "nonvolatile" ($c^* < 0.1 \ \mu g \ m^{-3}$), "semivolatile" (SVOC; 0.1 $\mu g \ m^{-3} < c^* < 1000 \ \mu g \ m^{-3}$), and "intermediate-volatility" (IVOC; 1000 $\mu g \ m^{-3} < c^* < 100,000 \ \mu g \ m^{-3}$) organic compounds. In other words, the *n*=8 basis set with 661 c* from 10^{-2} to $10^5 \ \mu g \ m^{-3}$. In this work, the *n*=8 set proves more than suitable for fitting 662 the existing α -pinene ozonolysis data, and for spanning simulated SOA partitioning data 663 across realistic ranges of temperature and organic aerosol concentration.

664 For modeling applications where a reduced n is desirable, (e.g. a global model) 665 the n=8 regression results can be translated to a reduced basis set, to polynomial fits, or to 666 lookup tables. For example, n=4 may be an attractive treatment for online partitioning calculations in a large-scale CTM. With n=4, the choice of c_{min}^* and c_{max}^* values will be 667 668 important in determining the performance of the partitioning fit, and use of equations 9 and 10 is recommended for basis vector selection. Figure 4 showed that n=4 basis 669 670 vectors were not acceptable, but the specific n=4 basis sets were not optimized for best performance in section 3.2. The basis vectors of $[0.1 \ 1 \ 10 \ 100]$, $[1 \ 10 \ 100 \ 10^3]$, and $[0.1 \ 10 \ 100]$ 671 672 2.2 46 1000] should be considered if an n=4 representation is desired. It should be 673 emphasized that an *n*=4 basis set may be rich enough for fitting the experimental data and 674 for the online partitioning calculations, but a basis vector with n>4 may be needed for 675 quantifying uncertainty limits during the data reduction step.

676

677 Summary and Conclusion

A method for fitting of smog chamber data for CTM applications was presented. The method uses a volatility basis set of compounds with fitted AMF and user-specified temperature-dependent saturation concentrations. A single effective enthalpy is fitted to match experimental and modeled temperature sensitivity of AMF. Important aspects of the method include: selection of minimum and maximum c* values, selection of the number of basis set compounds, and estimation of uncertainty bounds for the regressionprediction. Main conclusions include:

- When fitting parameters to experimental smog chamber data, the range of effective saturation concentration (c^*) values used must be consistent with the expected temperature and organic aerosol concentration range in the final model application. Equations for c^*_{min} and c^*_{max} are presented in section 2.2.
- The length of the basis vector (number of c* values used, n) can influence both the quality of the data fitting, and the ability to compute confidence intervals.
 Careful selection of c* values is required with n≤5 to insure a good fit and suitable confidence intervals. With 6 or more lognormally spaced c* values, the best fit line and confidence intervals become insensitive to the exact c* basis vector used.
- A method for determination of confidence intervals is demonstrated. It is useful to divide the confidence interval problem for AMF prediction from SOA chamber experiments into two different regimes: (1) confidence intervals for *interpolation within* experimental data; and (2) extrapolation *outside the data*. Interpolation CIs depend on the scatter in the experimental data itself.
- Although parts of the absorption SOA parameterization are nonlinear and/or
 implicit, the calculation of AMF is linear in the stoichiometric coefficients α_i.
 Therefore, fitted α_i values can be calculated using nonnegative linear least squares
 if the c_{OA,tot} and ΔH are fixed. This makes the regression computationally
 efficient.

In addition to traditional final AMF smog chamber experiments, newer dynamic aerosol formation experiments (e.g. PTRMS) and experiments with varied temperatures can be useful in constraining AMF parameterizations. Highly time-resolved data, such as a 1-min time series of AMF vs. ΔROG from a PTRMS and SMPS, may require averaging to reduce serial correlation between datapoints (discussed in section 3.5).

• Errors in apparent temperature dependence can be generated by correlated errors in measurements and temperature. This can be partially corrected with temperature ramp experiments that isolate the temperature sensitivity.

• Temperature-dependence of semivolatile aerosols is best dealt with by matching experimental and modeled temperature sensitivity ($\partial \ln c / \partial T$). For the α -pinene ozonolysis, the fitted temperature sensitivity at 40 µg m⁻³ and 20°C is -0.02 K⁻¹ and the associated ΔH is 33 kJ mol⁻¹. Fitted ΔH values depend on the basis vector (but temperature sensitivities do not); therefore, comparisons should be made of temperature sensitivity and not of ΔH .

In order to reduce uncertainty in CTM modeling, additional experimental data is
 needed at both lower concentrations *and* at a range of temperatures.

Although a large number value of n (~6 or more) may be required for an initial characterization of confidence intervals, this large number of parameters does not necessarily have to be used in the CTM. In other words, the items of value to the CTM modeler are the aerosol mass fraction and uncertainty bounds themselves, not the parameters, data, and error-correlation matrices used to generate them.

727 While each CTM modeling application may have a different optimal basis set, a • practical method of achieving improved CTM modeling results may for wider 728 reduction of experimental data to a basis set including "nonvolatile" ($c^* < 0.1 \text{ µg}$ 729 m⁻³), "semivolatile" (SVOC; 0.1 μ g m⁻³ < c* < 1000 μ g m⁻³), and "intermediate-730 volatility" (IVOC; 1000 μ g m⁻³ < C* < 100,000 μ g m⁻³) organic compounds. In 731 other words, the n=8 basis set with c* from 10^{-2} to $10^5 \ \mu g \ m^{-3}$. The parameters 732 733 can be reduced to a smaller set as needed for specific CTM applications as 734 needed.

735

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REFERENCES

743	An, W.J., Pathak, R.K., Lee, B.H. and Pandis, S.N., 2007. Aerosol volatility
744	measurement using an improved thermodenuder: Application to secondary
745	organic aerosol. Journal Of Aerosol Science, 38(3): 305-314.
746	Bian, F. and Bowman, F.M., 2002. Theoretical method for lumping multicomponent
747	secondary organic aerosol mixtures. Environmental Science & Technology,
748	36(11): 2491-2497.
749	Cocker, D.R., Clegg, S.L., Flagan, R.C. and Seinfeld, J.H., 2001. The effect of water on
750	gas-particle partitioning of secondary organic aerosol. Part I: alpha-pinene/ozone
751	system. Atmospheric Environment, 35(35): 6049-6072.
752	Donahue, N.M., Robinson, A.L., Stanier, C.O. and Pandis, S.N., 2006. Coupled
753	partitioning, dilution, and chemical aging of semivolatile organics. Environmental
754	Science & Technology, 40(8): 2635-2643.
755	Edgerton, E.S. et al., 1999. Particulate Air Pollution in Mexico City: A Col laborative
756	Research Project. Journal Of The Air & Waste Management Association, 49(10):
757	1221.
758	Griffin, R.J., Cocker, D.R., Flagan, R.C. and Seinfeld, J.H., 1999. Organic aerosol
759	formation from the oxidation of biogenic hydrocarbons. Journal Of Geophysical
760	Research-Atmospheres, 104(D3): 3555-3567.
761	Hamilton, L.C., 1992. Regression with Graphics. Wadsworth, Inc., Belmont, CA.
762	Hatakeyama, S., Izumi, K., Fukuyama, T. and Akimoto, H., 1989. Reactions Of Ozone
763	With Alpha-Pinene And Beta-Pinene In Air - Yields Of Gaseous And Particulate
764	Products. Journal Of Geophysical Research-Atmospheres, 94(D10): 13013-
765	13024.
766	Kamens, R.M. and Jaoui, M., 2001. Modeling aerosol formation from alpha-pinene plus
767	NOx in the presence of natural sunlight using gas-phase kinetics and gas-particle
768	partitioning theory. Environmental Science & Technology, 35(7): 1394-1405.
769	Kanakidou, M. et al., 2005. Organic aerosol and global climate modelling: a review.
770	Atmospheric Chemistry And Physics, 5: 1053-1123.
771	Kroll, J.H., Chan, A.W.H., Nga, N.L., Flagan, R.C. and Seinfeld, J.H., 2007. Reactions of
772	Semivolatile Organics and Their Effect on Secondary Organic Aerosol Formation.
773	Environmental Science & Technology, 41(10): 3545-3550.
774	Lee, A. et al., 2006. Gas-phase products and secondary aerosol yields from the
115	photooxidation of 16 different terpenes. Journal Of Geophysical Research-
//6	Atmospheres, $\Pi \Pi (D I /)$.
///	Lurmann, F. W. et al., 1997. Modelling urban and regional aerosols.2. Application to
778	California's South Coast Air Basin. Atmospheric Environment, 31(17): 2695-
790	2/10.
/80	Mathworks, 2002. MATLAB.
/81	Ng, N.L. et al., 2006. Contribution of first-versus second-generation products to
/82 782	secondary organic aerosols formed in the oxidation of diogenic hydrocarbons. Environmental Science & Technology, $40(7)$, 2292, 2207
183	Environmental Science & rechnology, $40(7)$: 2283-2297.
/84 795	Environmental Science & Technology 20(9): 2590 2595
183	Environmental Science & Technology, $30(8)$: 2380-2383.

786	Odum, J.R., Jungkamp, T.P.W., Griffin, R.J., Flagan, R.C. and Seinfeld, J.H., 1997. The
787	atmospheric aerosol-forming potential of whole gasoline vapor. Science,
788	276(5309): 96-99.
789	Offenberg, J.H., Kleindienst, T.E., Jaoui, M., Lewandowski, M. and Edney, E.O., 2006.
790	Thermal properties of secondary organic aerosols. Geophysical Research Letters,
791	33, L03816, doi:10.1029/2005GL024623.
792	Pandis, S.N., Harley, R.A., Cass, G.R. and Seinfeld, J.H., 1992. Secondary Organic
793	Aerosol Formation And Transport. Atmospheric Environment Part A-General
794	Topics, 26(13): 2269-2282.
795	Pankow, J.F., 1994a. An Absorption-Model Of Gas-Particle Partitioning Of Organic-
796	Compounds In The Atmosphere. Atmospheric Environment, 28(2): 185-188.
797	Pankow, J.F., 1994b. An Absorption-Model Of The Gas Aerosol Partitioning Involved In
798	The Formation Of Secondary Organic Aerosol. Atmospheric Environment, 28(2):
799	189-193.
800	Pathak, R.K. et al., in press, 2007. Ozonolysis of α-pinene: parameterization of secondary
801	organic aerosol mass fraction. Atmospheric Chemistry And Physics.
802	Pathak, R.K., Stanier, C.O., Donahue, N.M. and Pandis, S.N., 2007. Ozonolysis of alpha-
803	pinene at atmospherically relevant concentrations: Temperature dependence of
804	aerosol mass fractions (yields). Journal Of Geophysical Research-Atmospheres,
805	112(D03201, doi:10.1029/2006JD007436).
806	Philippin, S., Wiedensohler, A. and Stratmann, F., 2004. Measurements of non-volatile
807	fractions of pollution aerosols with an eight-tube volatility tandem differential
808	mobility analyzer (VTDMA-8). Journal Of Aerosol Science, 35(2): 185-203.
809	Presto, A.A. and Donahue, N.M., 2006. Investigation of alpha-pinene plus ozone
810	secondary organic aerosol formation at low total aerosol mass. Environmental
811	Science & Technology, 40(11): 3536-3543.
812	Presto, A.A., Hartz, K.E.H. and Donahue, N.M., 2005. Secondary organic aerosol
813	production from terpene ozonolysis. 2. Effect of NOx concentration.
814	Environmental Science & Technology, 39(18): 7046-7054.
815	Pun, B.K. et al., 2003. Uncertainties in modeling secondary organic aerosols: Three-
816	dimensional modeling studies in Nashville/Western Tennessee. Environmental
817	Science & Technology, 37(16): 3647-3661.
818	Robinson, A.L. et al., 2007. Rethinking organic aerosols: Semivolatile emissions and
819	photochemical aging. Science, 315(5816): 1259-1262.
820	Salcedo, D. et al., 2006. Characterization of ambient aerosols in Mexico City during the
821	MCMA-2003 campaign with Aerosol Mass Spectrometry: results from the
822	CENICA Supersite. Atmospheric Chemistry And Physics, 6: 925–946.
823	Seber, G.A.F. and Wild, C.J., 2003. Nonlinear Regression. John Wiley and Sons.
824	Seinfeld, J.H. and Pankow, J.F., 2003. Organic atmospheric particulate material. Annu.
825	Rev. Phys. Chem., 54: 121-140.
826	Stanier, C.O., Patnak, R.K. and Pandis, S.N., 2007. Measurements of the volatility of
827	Aerosols from alpha-Pinene Ozonolysis. Environmental Science and Technology,
828	41(8): 2/30-2/03.
829	Strader, K., Lurmann, F. and Pandis, S.N., 1999. Evaluation of secondary organic aerosol
830	iormation in winter. Atmospheric Environment, 33(29): 4849-4863.

- Takekawa, H., Minoura, H. and Yamazaki, S., 2003. Temperature dependence of
 secondary organic aerosol formation by photo-oxidation of hydrocarbons.
 Atmospheric Environment, 37(24): 3413-3424.
- Tsigaridis, K. and Kanakidou, M., 2003. Global modelling of secondary organic aerosol
 in the troposphere: a sensitivity analysis. Atmospheric Chemistry And Physics, 3:
 1849-1869.
- Yu, J.Z., Cocker, D.R., Griffin, R.J., Flagan, R.C. and Seinfeld, J.H., 1999. Gas-phase
 ozone oxidation of monoterpenes: Gaseous and particulate products. Journal Of
 Atmospheric Chemistry, 34(2): 207-258.
- 840
- 841

0.1	I WOIC I MINIM	iaini ana m		, aia							
	Scenario	c _{OA,max}	δ	f	T_{min} - T_{max}	$c*_{min}(25^{\circ}C)$	$c^*_{max}(25^\circ C)$	recommended			
		(µg m ⁻³)	(µg m ⁻³)		(°C)	(µg m ⁻³)	(µg m ⁻³)	n			
	General	60	0.1	0.05	0 - 40	1×10^{-3}	240,000	5-10			
	Remote with high	5	0.05	0.05	-30 - 40	$7x10^{-4}$	4.5×10^{6}	6-11			
	altitude										
	Polluted urban	120	0.5	0.05	0 - 40	0.07	480,000	4-8			
	Proposed decadal	25	0.07	0.05	0 - 40	0.01	100,000	4-8			
0.45	range										
845											
846											
847											
010	Table) Error	Formulas	Ugad for ("aadma	an of Eit						
040		Formulas		Jooune							
	Name & Symbol				Formul	Formula					
			D /		1 /	1 ()					
	Mean Squared Erro	or (err _{mse}) vs	. Data		$\frac{1}{2}\sum \left(\mathcal{E}_{1} \right)$	$\frac{1}{2}\sum \left(\varepsilon_{\perp}-\hat{\varepsilon}_{\perp}\right)^2$					
					m^{-1}	ji)					
	Average Bias vs. [Data (err _{man})			1 - (\sim					
	Average bias vs. Data (einmean) $\frac{1}{-\sum_{i} (\xi_i - \hat{\xi})}$										
	m — (······)										
	Mean Absolute Fra	actional Erro	r vs. Data (e	. ()=	$\left(\begin{vmatrix} z & \hat{z} \end{vmatrix} \right)$						
			× ×	nuc)	$\frac{1}{\sum} \left \frac{ \varsigma_i }{ \varsigma_i } \right $	$\frac{1}{\sum} \left \frac{ \varsigma_i - \varsigma }{ \varsigma_i - \varsigma } \right $					
					$m \sim$	ξi					

Table 1 Minimum and maximum c* values for some scenarios

List of data sources for a-pinene SOA AMF Table 3

Study	Experiments Used for Fitting in This Work	α-pinene Concentration Range	Temperature Range	No. Experiments (for this analysis)
Griffin et al., 1999	Outdoor (dark) chamber reaction with O ₃ in excess, OH scavenger, inorganic seed, and dry conditions	15-65 ppb	32-37 °C	6
Hoffmann et al., 1997	Outdoor (dark) chamber reaction with O_3 in excess and OH (no OH scavenger), inorganic seed, and dry conditions	88-154 ppb	46-50 °C	5
Cocker et al., 2001	Indoor chamber reaction with O3 in excess or limiting, OH scavenger, RH < 20%; some seeded experiments and some unseeded	23-163 ppb	28-30 °C	15
Pathak et al., 2007	Indoor chamber reaction with O ₃ in excess, OH scavenger. Seeded and unseeded experiments at a many temperature / concentration combinations	4-50 ppb	0-40 °C	31
Presto and Donahue, 2006	Indoor chamber reaction with O_3 in excess, low NOx. Experiments $6/14/05$ and $6/28/05$.	1.5-138 ppb	22 °C	2

Table 4. Description of synthetic data experiments

Q	5	7
0	J	1

855

Chemical and	Pseudodata Sets					
	Generated from Each					
Description		1-5		6-10	11-14	Underlying Model
2 SOA Products	log ₁₀ c*	0.699; 2.90				Pseudodata A, B1, B2
	α x100	10; 15				(see descriptions
	$\Delta H (kJ mol^{-1})$	80; 60				below)
14 504	log ₁₀ c*	-1.75; -1.54; -0.43	6; -0.218; +0.60	0.702; 1.04; 1.56; 1.84; 2.05	2.58; 3.02; 3.24; 3.83	
products	α x100	0.12; 0.06; 0.6;	3; 1.8	4.2; 1.2; 0.9; 2.4; 0.6	9.0; 1.2; 0.6; 2.4	Pseudodata C
	$\Delta H(kJ mol^{-1})$	120; 110; 90;	80; 60	45; 72; 45; 83; 100	40; 35; 40; 50	

9 Pseudodata:

A. Hypothetical sampling by 19 chamber experiments. Very low sampling error and no bias. Experimental c_{OA} values from 7-525 μg m⁻³ and temperatures from 26 to 37°C.

B1. Hypothetical sampling by 3 series of chamber experiments. Series one: 10 experiments at 26°C with low error but 10% AMF bias high. Series 2: 6 experiments at 37°C with low error but 10% AMF bias low. Series 3: 5 experiments with constant reactant concentration, but with different temperature in each case (17-37°C); low error; no bias.

863
 B2. Same as B1 but adding a 4th series: 5 temperature ramp experiments so that 5 AMF measurements are taken during each experiment, with temperatures ranging from 22-39°C.

865 C. Hypothetical sampling by 19 chamber experiments. Very low sampling error and no bias. Experimental c_{OA} values from 7-525 µg m⁻³ and temperatures from 26 to 37°C. (Identical to Pseudodata A).

Appendix A –AMF (yield) equations with density and preexisting organic aerosol considerations

869

870 Equation 6 is the commonly encountered fitting equation. Here we show that equation 6 is a

871 special case of a more general set of mass balances for the semivolatile aerosol system. The

equations in this appendix are necessary when accounting for experimental and simulation cases

873 where the organic aerosol density is not assumed to be unity, and for cases with preexisting

874 organic mass.

876
$$c_{i,gas} \equiv y_i c_i^* = \frac{c_{aeri,i}}{c_{OA}} c_i^*$$
(A1)

the following expression can be derived by mass balance:

878
$$\xi_{i} = \frac{c_{aeri,i}}{c_{tot,i}} = \frac{1}{1 + \frac{c_{i}^{*}}{c_{OA}}}$$
(A2)

879 This can be used with the following definitions of AMF and reaction mass yield α_i .

$$\Delta c_{tot,i} \equiv \alpha_i \Delta ROG \tag{A3}$$

881
$$AMF = \xi \equiv \frac{\Delta c_{OA}}{\Delta ROG} = \sum \frac{\alpha_i}{1 + \frac{c_i^*}{c_{OA}}}$$
(A4)

Using the above equations, we can write expressions for prediction of organic aerosol mass (c_{OA}), species-specific partitioning values ε_i , and smog chamber AMF as function of parameters such as α_i , Δ ROG, and c_i^* . To keep the expressions general, we assume the density of the organic aerosol is unknown, and with preexisting organic aerosol.

We denote the values of variables corresponding to the real organic aerosol density as in the above equations (no subscript). We denote values of variables under the case of assumed unit organic aerosol density with a subscript $\rho=1$ for physical variables such as ξ and c_{OA} . Preexisting aerosol is denoted by the superscript P.

891 Expression for
$$c_{OA}$$
, $c_{OA,\rho=1}$, ξ and $\xi_{\rho=1}$ are then:

$$892 c_{OA} = c_{OA,\rho=1} \frac{\rho}{\rho_1} = \sum \xi_i c_{i,tot} = \sum \frac{c_{tot,i}}{1 + \frac{c_i^*}{c_{OA}}} = \sum \frac{c_{tot,i}}{1 + \frac{c_i^* \rho_1}{c_{OA,\rho=1} \rho}} = \sum \frac{c_{tot,i}}{1 + \frac{c_i^*}{c_{OA,\rho=1}}} = \sum \frac{c_{tot,i}}{1 + \frac{c_i^* \rho_1}{c_{OA,\rho=1}}}$$
(A5)

where the definitions $c_i^{*'} = \frac{c_i^* \rho_1}{\rho}$ and $\alpha_i^{'} = \frac{\alpha_i \rho_1}{\rho}$ are used. The 4 alternate versions of the denominator in equation A5 can be made in all the cases below, and the different expressions are not repeated. Separating the organic aerosol (c_{OA}) and the total concentration of species i into preexisting and reaction-generated parts, gives the equations for c_{OA} and AMF in the case of preexisting aerosol and assumed unit density. An equation for AMF in the case of preexisting aerosol is not written, since the organic aerosol concentration can change from partitioning of preexisting and reaction-generated species.

900
$$c_{OA} = \sum \xi_{i} c_{i,tot} = \sum \frac{c_{tot,i}^{P} + \Delta c_{tot,i}}{1 + \frac{c_{i}^{*}}{c_{OA}}} = \sum \frac{c_{tot,i}^{P} + \alpha_{i} \Delta ROG}{1 + \frac{c_{i}^{*}}{c_{OA,\rho=1}} \rho / \rho_{1}} = \sum \frac{c_{tot,i}^{P} + \alpha_{i}' \rho / \Delta ROG}{1 + \frac{c_{i}^{*} \rho / \rho_{1}}{c_{OA}}}$$
(A6)

901 In the case of no preexisting aerosol:

902
$$\frac{c_{OA,\rho=1}}{\Delta ROG} = \xi_{\rho=1} = \frac{\rho_1}{\rho} \sum \frac{\alpha_i}{1 + \frac{c_i^* \rho_1}{c_{OA,\rho=1} \rho}} = \sum \frac{\alpha_i'}{1 + \frac{c_i'}{c_{OA,\rho=1}}}$$
(A7)

903 and
$$\xi = \frac{\rho}{\rho_1} \xi_{\rho=1} = \frac{\rho}{\rho_1} \sum \frac{\alpha'_i}{1 + \frac{c_i^{*'}}{c_{OA,\rho=1}}} = \frac{\rho}{\rho_1} \sum \frac{\alpha'_i}{1 + \frac{c_i^{*'} \rho}{c_{OA} \rho_1}}$$
(A8)

904 and
$$c_{OA} = \Delta ROG \frac{\rho}{\rho_1} \sum \frac{\alpha'_i}{1 + \frac{c_i^* \rho}{c_{OA} \rho_1}}$$
(A9)

In this work, all aerosol concentrations and AMFs are calculated using unit density conversions from SMPS aerosol volume, so the fitting really recovers the normalized values $c_i^{*} = \frac{c_i^* \rho_1}{\rho}$ and $\alpha_i^{'} = \frac{\alpha_i \rho_1}{\rho}$ according to equation A7. To make concentration predictions for organic aerosols with densities other than 1.0 requires the use of equation A9. At low aerosol mass fractions, the terms cancel and correct AMF and c_{OA} predictions will be made by regressed values of α , regardless of the density assumption made. At high aerosol loadings, AMF and c_{OA} concentration predictions will be low for actual aerosol specific gravities > 1.

913 Appendix B – Derivation of formulae for c*_{min} and c*_{max}

914

915 For any expected parameterization applicability domain (defined by $c_{OA,min}$, $c_{OA,max}$, T_{min} , 916 and T_{max}) the extreme values required of the effective saturation concentration (c_{min} and c_{max}) 917 are calculated as follows. For a single condensing component in the presence of $c_{NV,min}$ 918 (nonvolatile, organic, and solution forming) preexisting aerosol, the mass balance *for the* 919 *semivolatile species* is:

$$c_{aer,SV} = c_{tot,SV} - c_{gas,SV} = c_{tot,SV} - y_{SV}c_{SV}^*$$
(B1)

921 where $c_{gas,SV}$ and $c_{aer,SV}$ refer to the gas- and particle-phase concentrations of the semivolatile in 922 $\mu g m^{-3}$. Adding in the preexisting nonvolatile organic, the total organic aerosol concentration is:

923
$$c_{OA} = c_{aer,SV} + c_{NV,\min} = c_{tot,SV} - y_{SV}c_{SV}^* + c_{NV,\min}$$
(B2)

To determine c_{\min}^* , the lowest c^* necessary to achieve the required modeling accuracy goal δ , we construct a scenario where a truly non-volatile species is represented in the model as having a finite saturation concentration c_{\min}^* , therefore creating an error. The true aerosol concentration in this case will be:

928

$$c_{OA,true} = c_{tot,SV} + c_{NV,\min}$$
(B3)

929 while the modeled aerosol concentration will be:

930 $c_{error} = y_{SV} c_{\min}^* \le \delta$ (B4)

931 to keep this error less than a user-selected goal δ :

932 $c_{\min}^* \le \frac{\delta}{y_{SV}}$ (B5)

933 If the organic aerosol consists only of the nonvolatile and semivolatile fractions, then we can
934 eliminate y_{SV} to give:

935
$$c_{\min}^* \le \delta \left(1 + \frac{c_{NV,\min}}{c_{tot,SV}} \right)$$
(B6)

For a conservative estimate of c_{min}^* , we assume $c_{tot,SV} >> c_{NV,min}$, to keep the right hand side as small as possible. Furthermore, the calculation is assumed to apply at the highest temperature applied in the model (T_{max}).

939
$$c_{\min}^{*}(T_{ref}) = \delta \frac{T_{\max}}{T_{ref}} \exp\left[\frac{\Delta H}{R} \left(\frac{1}{T_{\max}} - \frac{1}{T_{ref}}\right)\right]$$
(B7)

940 Putting in 100 kJ/mol, neglecting T_{max}/T_{ref} , and with temperature in Kelvin:

941
$$c_{\min,298K}^* \approx \delta \exp\left[\frac{12000}{T_{\max}} - 40.27\right]$$
 (B8)

942 With δ of 1 μg m $^{-3}$ and T_{max} at 40°C, $c*_{min}$ is 0.14 μg m $^{-3}.$

943 To derive c*_{max} we consider a case with non-volatile solution-forming aerosol (secondary or 944 primary) $c_{NV,max}$ and one condensing species with saturation concentration c_{max}^* . The goal is 945 then to select c_{max}^* sufficiently high so that material in this bin is not in the aerosol phase. 946 Donahue, Robinson et. al (2006) suggest 1000 $c_{NV,max}$ which ensures that less than about 0.1% of 947 material in the c_{max}^* bin partitions to the aerosol phase. A slightly more flexible approach is to 948 include the CTM modeler's tolerance for error in semivolatile partitioning under high 949 concentrations and low temperature. The mass balance for the semivolatile species is given by B1. A relative accuracy goal, denoted by f, is defined. At the high end of the expected 950 concentration range in the model, the accuracy goal ($\mu g m^{-3}$) is $\pm f c_{oa,max}$ (e.g., f=0.1 for $\pm 10\%$ 951 952 accuracy).

953
$$c_{aer,SV} \le f(c_{NV,\max} + c_{aer,SV})$$
(B9)

954 If we assume that $c_{aer,SV}$ is a small fraction of $c_{NV,max}$ then:

955
$$c_{aer,SV} = \frac{c_{tot,SV}}{1 + \frac{c_{max}^*}{c_{NV,max}}} \le fc_{NV,max}$$
(B10)

956 which can be solved for c^*_{max} yielding:

957
$$c_{\max}^* \ge \frac{c_{tot,SV}}{f} - c_{NV,\max}$$
(B11)

958 The $c_{tot,SV}$ term is difficult to estimate. It represents the total pool of semivolatile compounds 959 that need to be modeled as in the gas phase. One reasonable assumption is to assume for any 960 given location, that gas phase semivolatile concentrations are proportional to organic aerosol (or 961 vis versa), such that $c_{tot,SV} = kc_{NV,max}$. This simplifies B11 to:

962
$$c_{\max}^* \ge c_{NV,high} \left(\frac{k}{f} - 1\right)$$
(B12)

963 If we further assume that k is at least unity (the mass of semivolatiles is at least the organic964 aerosol concentration) and the desired fractional error is small, then:

965
$$c_{\max}^* \ge \frac{k}{f} c_{NV,high}$$
(B12)

where k is an unknown ratio of the semivolatiles gases to the organic aerosol. Taking values from Mexico City, we find a total VOC concentration of ~2000 μ g m⁻³ (Edgerton et al., 1999) and an organic aerosol concentration of ~20 μ g m⁻³ (Salcedo et al., 2006). If 5% of the VOC is SV, then k = 5 for Mexico City. Applying a temperature correction similar to the one for c*_{min} (but in the other direction)

971
$$c_{\max}^{*}(T_{ref}) \ge \frac{k}{f} c_{NV,\max} \frac{T_{\min}}{T_{ref}} \exp\left[\frac{\Delta H}{R} \left(\frac{1}{T_{\min}} - \frac{1}{T_{ref}}\right)\right]$$
(B13)

972 Using T_{ref} of 25 °C, ΔH of 100 kJ mol⁻¹, k=5, and ignoring T_{min}/T_{ref} , equation B13 reduces to:

973
$$c_{\max,298K}^* \approx \frac{5}{f} c_{NV,\max} \exp\left[\frac{12000}{T_{\min}} - 40.27\right]$$
 (B14)

974 If the fractional accuracy goal is 10%, $c_{NV,max}$ is 100 µg m⁻³, and T_{min} is 0°C then $c*_{max}(298 \text{ K}) \approx$

- 975 200,000 μ g m⁻³, or 200 mg m⁻³. Thus the total multiplier to c_{NV,max} in this case is 2000,
- 976 comparable to the multiplier of 1000 in Donahue et al. (2006).

978 1. Create a vector $\mathbf{c}_{OA,model}$ (length $p \sim 20$) of aerosol concentrations, logarithmically spaced from $c_{OA,min}$ to $c_{OA,max}$ (the maximum and minimum expected aerosol concentrations in 979 980 the modeling application). 2. Create a vector (length j) of temperatures from T_{min} to T_{max} . In this work, j is selected for 981 982 spacing between elements of 10°C. 3. Let the experimental yields be a vector ξ , and the fitted predictions at the experimental 983 values of c_{OA} and T are the vector $\hat{\boldsymbol{\xi}} \pm \boldsymbol{\delta}$ where the values $\boldsymbol{\delta}$ are from the confidence 984 985 intervals of the asymptotic method (e.g., MATLAB nlpredci function). All these vectors $(\xi, \hat{\xi}, \text{ and } \delta)$ will have length of *m* corresponding to number of experiments. 986 4. The experimental data are shifted by δ such that: $\xi_{high} = \xi + \delta$ and $\xi_{low} = \xi - \delta$. 987 988 5. A reference error value, corresponding to how well the fit goes through the data, is noted: $MSE_{tef} = \sum \left(\xi_i - \hat{\xi}_i\right)^2 / m$ 989 990 6. Initialize *j* upper confidence intervals (CI) vectors (each of length p), one for each 991 temperature. Set each element to zero. 992 7. Initialize *i* lower confidence interval vectors (each of length *p*), one for each temperature. 993 Set each element to a large value (e.g. 999). 994 8. Create a vector (length k) of ΔH values. In this work, k was usually selected for a separation of 5 kJ mol⁻¹ and a range from 10 to 120 kJ mol⁻¹. 995 996 The outer loop (steps 9-16) is repeated $4 \cdot k$ times. For each value of ΔH , the four runs are 997 for upper CI (allowing random selection of low volatility yields); lower CI (allowing 998 random selection of low volatility yields); upper CI (allowing random selection of high 999 volatility yields); and lower CI (allowing random selection of high volatility yields).

977 Appendix C. Algorithm for the Monte-Carlo confidence intervals

1000

OUTER LOOP

- 10019. Since for any iteration ΔH is fixed, the left hand matrix of equation 13 (C) is calculated1002and stored. This speeds up the subsequent calculation of yields given any vector $\boldsymbol{\alpha}$.
- 1003 10. The *n* yields $\boldsymbol{\alpha}$ are divided into two groups, those that will be selected randomly, and 1004 those that will be fit. If the iteration involves randomized low volatility yields, then some 1005 number of the lowest volatility species (e.g. α_1 and α_2) are selected randomly. If the 1006 iteration involves randomized high volatility yields, then some of the highest volatility 1007 species (e.g., α_{n-1} and α_n) are chosen randomly.
- 1008 In this work, the number of low and high volatility yields to be randomized were selected 1009 by the user on a case by case basis, with the goal of only selecting yields that are poorly 1010 constrained by the experimental data. The typical n=7 case required randomization of α_{l_1} , α_2 and α_6 , α_7 . Inspection of the error covariance matrix and the standard errors on the 1011 1012 original fitted yields will help in selecting which yields to randomize. In the 1013 demonstration case of Figure 3 (section 3.1), the α 's with the 5 largest standard errors 1014 were selected, or elements 1, 2 (uncertain low volatility compounds) and 6, 7 and 8 1015 (uncertain high volatility compounds). The range of experimental data can serve as a guide for which terms are poorly constrained. If a given c* value is outside the range of 1016 experimental data (in this case from 7-525 μ g m⁻³), then the corresponding α value should 1017 1018 probably be randomized. Erring on the conservative side, and trying to select α values 1019 randomly that are in fact well constrained by the data does not effect the quality of the 1020 result. Rather it adds run time as calculations are performed that are of no subsequent use 1021 in determining the CIs.

1022 11. Let the number of random α values be q and the number to be fit equals n-q. Let the 1023 fixed yields be denoted as α_q and the variable yields as α_{n-q} . Also let the C matrix (from 1024 equation 13) be split into an $m \ge q$ matrix C_q corresponding to the columns of the fixed 1025 yields, and a matrix C_{n-q} [size $m \ge (n-q)$] corresponding to the columns of the adjustable 1026 yields.

1027 INNER LOOP

1028 12. 1000 random combinations of the random α values are chosen.

1029 13. For each of the 1000 cases, the remaining n-q yields are calculated by solving a 1030 modified version of equation 13. The adjustable yields a_{n-q} are determined by 1031 solving $C_{n-q}a_{n-q} = \xi^* - C_qa_q$ with $a_{n-q} \ge 0$ where ξ^* is ξ_{high} when calculating an 1032 upper CI and ξ^* is ξ_{low} when calculating a lower CI.

1033 14. For each of the 1000 cases, the mean squared error is calculated 1034 $MSE = \sum \left(\xi_i - \hat{\xi}_i (\boldsymbol{a_q}, \boldsymbol{a_{n-q}})^2 / m\right).$

1035 If $MSE \leq MSE_{ref}$ then it is a suitably good fit. Fitted values $\hat{\xi}$ are calculated at all *j* x 1036 *p* combinations of temperature and organic aerosol concentration. For upper CI, 1037 $CI_{high} = max(CI_{high}, \hat{\xi})$, and for the lower CI, $CI_{low} = min(CI_{low}, \hat{\xi})$ where the 1038 comparison is on an element by element basis.

- 1039 15. If the confidence intervals moved more than a present tolerance, repeat the inner loop by1040 going to step 12.
- 1041 16. Loop back to step 9 until all 4*k* cases are completed.

1042 17. For each temperature of interest, upper and lower confidence interval vectors are now
1043 calculated using the Monte-Carlo method. They should be inspected graphically and
1044 compared to the limits from the asymptotic method.

1047 Supplemental Data Table (online only)

Author's	ΔROG	ΔROG	O ₃	Temperature	RH	OH-	Seed	-	Result
Identifier	(ppb)	(µg	initial	(°C)		Scavenger		ΔΜ	AMF
		m^{-3}	(ppb)			C		(µg	$\Delta M / \Delta ROG$
								m ⁻³)	(%)
Single-temp	berature fi	nal AMF	chambe	r experiments			•		
Griffin et al	., 1999								
6/5/98a	16.7	89		36.8		2-butanol	$(NH_4)_2SO_4$	7.4	8.3
6/5/98b	18.2	97		36.8		2-butanol	$(NH_4)_2SO_4$	8.5	8.7
6/7/98a	31.0	169		30.2		2-butanol	$(NH_4)_2SO_4$	30.3	17.9
6/7/98b	45.5	249		30.2		2-butanol	$(NH_4)_2SO_4$	46.0	18.4
6/9/98a	57.0	307		34.8		2-butanol	$(NH_4)_2SO_4$	52.3	17.0
6/9/98b	65.0	350		34.8		2-butanol	$(NH_4)_2SO_4$	65.1	18.6
Hoffmann e	et al., 199	7							
9/27A	88.0	454	210	48.1		none	$(NH_4)_2SO_4$	82.0	18.0
9/27B	38.0	196	220	48.9		none	$(NH_4)_2SO_4$	29.9	15.2
9/29A	101.9	795	320	47.2		none	$(NH_4)_2SO_4$	80.4	23.0
9/29B	154.1	527	327	48.1		none	$(NH_4)_2SO_4$	183.5	15.2
10/02A	132.9	690	210	46.1		none	$(NH_4)_2SO_4$	94.9	13.7
Cocker et a	1., 2001				1 .				
11/29a	74.7	411	280	28.0	<2	2-butanol	$(NH_4)_2SO_4$	82.0	20.0
11/29b	144.1	792	360	28.0	<2	2-butanol	$(NH_4)_2SO_4$	190.0	24.0
12/03a	40.6	223	370	28.0	<2	2-butanol	$(NH_4)_2SO_4$	34.0	15.2
12/03b	72.3	398	315	28.0	<2	2-butanol	$(NH_4)_2SO_4$	79.0	19.8
12/06b	22.9	126	280	28.0	<2	2-butanol	$(NH_4)_2SO_4$	10.0	7.9
01/08a	72.4	399	270	27.8	<2	2-butanol	none	80.0	20.1
01/10a	41.0	225	278	28.9	<2	2-butanol	none	35.0	15.5
01/10a	93.4	512	499	28.9	<2	2-butanol	none	110.0	21.5
01/13a	72.3	396	244	29.5	14.1	2-butanol	$(NH_4)_2SO_4$	79.0	19.9
01/136	118.4	648	352	29.5	14.4	2-butanol	$(NH_4)_2SO_4$	145.0	22.4
01/14a	48.4	265	402	29.6	14.6	2-butanol	$(NH_4)_2SO_4$	48.0	18.1
01/14b	98.8	541	313	29.6	16.6	2-butanol	$(NH_4)_2SO_4$	115.0	21.3
01/18a	116.1	637	130	29.1	<2	2-butanol	$(NH_4)_2SO_4$	140.0	22.0
01/22a	162.3	888	175	29.6	<2	2-butanol	none	220.0	24.8
01/22b	139.3	763	380	29.6	<2	2-butanol	none	185.0	24.3
Pathak et al	., 2007		0.50	40.0	.10	0.1 1		7 0	
	14.3	76	250	40.0	<10	2-butanol	none	7.0	9.2
2	42.0	242	250	15.0	<10	2-butanol	none	50.7	21.0
3	38.3	217	250	20.0	<10	2-butanol	none	33.7	15.6
4	50.0	273	250	30.0	<10	2-butanol	none	48.7	17.8
5	50.0	265	250	40.0	<10	2-butanol	none	34.7	13.1
/	7.3	41	1500	20.0	<10	2-butanol	none	2.8	6.8
9	3.7	21	3100	20.0	<10	2-butanol	none	0.7	3.3
10	7.3	39	1500	40.0	<10	2-butanol	none	2.1	5.4
11	7.3	42	1500	15.0	<10	2-butanol	none	2.6	6.2
12	/.3	40	1500	30.0	<10	2-butanol	none	2.5	6.3
13	38.3	217	250	20.0	<10	2-butanol	$(NH_4)_2SO_4$	37.6	17.4
14	1.3	41	1500	20.0	<10	2-butanol	$(NH_4)_2SO_4$	2.9	0./

Author's	ΔROG	ΔROG	O ₃	Temperature	RH	OH-	Seed		Result
Identifier	(ppb)	(µg	initial	(°C)		Scavenger		ΔM	AMF
		m ⁻³)	(ppb)			_		(µg	$\Delta M/\Delta ROG$
								m ⁻³)	(%)
15	8.0	45	250	20.0	<10	2-butanol	$(NH_4)_2SO_4$	2.6	5.7
19	8.0	46	1500	15.0	<10	2-butanol	$(NH_4)_2SO_4$	3.7	8.0
20	7.3	39	1500	40.0	<10	2-butanol	$(NH_4)_2SO_4$	1.4	3.6
21	8.5	46	1500	30.0	<10	2-butanol	$(NH_4)_2SO_4$	3.4	7.3
22	14.3	82	250	15.0	<10	2-butanol	$(NH_4)_2SO_4$	7.2	8.8
23	14.3	82	300	15.0	<10	2-butanol	$(NH_4)_2SO_4$	9.3	11.3
24	14.3	81	250	20.0	<10	2-butanol	none	9.8	12.1
25	14.3	78	250	30.0	<10	2-butanol	none	10.0	12.8
26	38.3	220	250	15.0	<10	2-butanol	$(NH_4)_2SO_4$	33.3	15.1
27	38.3	203	250	40.0	<10	2-butanol	$(NH_4)_2SO_4$	37.0	18.3
28	29.0	167	250	15.0	<10	2-butanol	$(NH_4)_2SO_4$	18.2	10.9
29	7.3	41	750	20.0	<10	2-butanol	$(NH_4)_2SO_4$	3.1	7.5
30	33.0	180	250	30.0	<10	2-butanol	$(NH_4)_2SO_4$	26.5	14.7
31	14.3	78	250	30.0	<10	2-butanol	$(NH_4)_2SO_4$	7.6	9.7
32	15.0	91	250	0.0	<10	2-butanol	$(NH_4)_2SO_4$	16.1	17.7
33	17.0	98	250	15.0	<10	2-butanol	$(NH_4)_2SO_4$	10.5	10.7
34	14.3	78	250	30.0	<10	2-butanol	$(NH_4)_2SO_4$	10.8	13.8
35	24.0	127	250	40.0	<10	2-butanol	$(NH_4)_2SO_4$	16.0	12.6
36	42.0	255	250	0.0	<10	2-butanol	$(NH_4)_2SO_4$	96.2	37.7
PIRMS Ex	periments	3000							
Presto and	Donahue,	2006	200	22.0					
6/14/2005	1.50	0.4	280	22.0		none	none	0.001	0.017
	1.50	8.4						0.001	0.017
	2.47	15.9						0.037	0.41
	2.07	22.5						0.313	1.94
	4.01	22.3						0./1/	5.18
	3.93	22.2						1.12	5.00
	4.43	25.0						1.43	6 37
	4.78	20.9						1.71	6.99
	5.52	31.0						2.07	6.69
	5.17	29.0						2.07	7.63
	5 49	30.8						2.21	7.53
	5.85	32.8						2.32	7.30
	5.69	31.9						2.10	7.38
	6.04	33.9						2.56	7.76
	6.00	33.7						2.50	7.30
	6.06	34.0						2.63	7.72
	6.18	34.7					<u> </u>	2.68	7.73
	6.22	35.0						2.70	7.74
	6.06	34.0						2.74	8.05
	6.23	35.0						2.72	7.79
	6.21	34.9						2.74	7.85
	6.06	34.0						2.76	8.10
	6.33	35.5						2.75	7.74

Author's	ΔROG	ΔROG	O ₃	Temperature	RH	OH-	Seed	Result	
Identifier	(ppb)	(µg	initial	(°C)		Scavenger		ΔM	AMF
		m^{-3})	(ppb)					(µg	$\Delta M / \Delta ROG$
								m ⁻³)	(%)
	6.28	35.3						2.73	7.75
	6.07	34.1						2.75	8.07
	6.25	35.1						2.73	7.79
	6.21	34.9						2.76	7.91
	6.46	36.3						2.74	7.56
	6.16	34.6						2.74	7.92
	6.25	35.1						2.74	7.82
	6.29	35.3						2.73	7.75
	6.31	35.4						2.73	7.72
	6.38	35.8						2.74	7.66
	6.30	35.4						2.71	7.66
	6.28	35.3						2.72	7.70
	6.20	34.8						2.71	7.80
	6.29	35.3						2.71	7.67
	6.29	35.3						2.71	7.66
6/28/2005			390	22.0		none	none		
	24.1	135						15.3	11.3
	50.2	282						46.3	16.4
	71.1	399						73.7	18.4
	84.7	476						95.1	20.0
	96.8	543						112.9	20.8
	113.1	635						143.7	22.6
	129.5	727						164.9	22.7
	128.8	723						167.6	23.2
	129.2	726						170.3	23.5
	129.6	728						174.0	23.9
	134.1	753						185.0	24.6
	138.1	775						191.1	24.7
	138.8	779						190.7	24.5
	138.9	780						192.7	24.7
	139.8	785						191.7	24.4