Chapter 2 Describing Turbulence: Averages, Correlations and Spectra

Systeral at temporal veloary & pressure fields for incompensible fur weart flow obey the waven - Stoken equition V.U = DUL =0 e = = e (= + U.0 2) = eg - 0p + NO2U  $e\left(\frac{\partial u_{i}}{\partial z} + u_{j}\frac{\partial u_{i}}{\partial x_{i}}\right) = eg_{i} - \frac{\partial P}{\partial x_{i}} + \frac{\partial^{2}u_{i}}{\partial x_{i}^{2}}$  $U(X, t) = (U_1, U_2, U_3) \quad X = (X_1, X_2, X_3)$ =(U, V, W) = (x, y, z)P = P(x, t)inertia = Soly force - pressure + Viscons due gout geolet fine diffusion meeting meludes U. V V wonlinear mechanistim I which perturbation may reff-remporce Visionen defpern deriver for molecular momentum deffection " for nondense gover due to molecular collision, whereas for lequids due to local inter molecular Estesson avising from her dose proximity

Sufficien Smooth flow perbutition, whereas nonlinear mertia may Strengthen them such that their effets are often in Conflict X= X/L U' = 2/0 M JUL = M U/L2 QUIDU: =QU2/L Marvie / Widow = MU2/RUZ/L = The Re= UL/V U=mls V= M/R X=m M = WS = My Small Re Viscous Affresion domination, where are for large q = hy/m3 ne mertia dominater, i.e., V = M2/S Re most imported parameter destinguish stability of transform to Surfalance

Recall viscous diffusion layers:

Viscous layers:

Sudden acceleration flat plate:  $u_t = v u_{yy}$   $\delta = 3.64 \sqrt{vt}$  u(y,0) = 0 u(0,t) = U $u(\infty,t) = 0$ 

Layer grows in time due viscous diffusion

Oscillating flat plate: 
$$u_t = \upsilon u_{yy}$$
  
 $\delta = 6.5\sqrt{v/\omega}$   
 $u(0,t) = U_0 \cos \omega t$   
 $u(\infty,t) = 0$ 

Layer confined constant thickness

Stagnation point flow:  $\delta = 2.4\sqrt{\nu/B}$  layer not a function of x since convection balances diffusion

Flat plate boundary layer:  $u_{x} + v_{y} = 0$   $uu_{x} + vu_{y} = Uu_{yy}$   $\delta = 4.9\sqrt{vx/U}$  u(x,0) = 0  $u(x,\infty) = U$ 

Layer grows with x due convection

Ekman Layer on Free Surface: effects due to wind shear  $\tau(x)$ 

$$\delta = \sqrt{\frac{2\nu}{f}} = \text{Ekman layer thickness}$$

Coriolis force =  $2\underline{\Omega} \times \underline{V} = -fv\dot{i} + fu\dot{j} - 2\Omega\cos\theta u\dot{k} f = 2\Omega\sin\theta$ = planetary vorticity = 2\* vertical component  $\Omega$   $-fv = vu_{ZZ} fu = vv_{ZZ}$   $\mu u_z = \tau$  at  $z = 0 \ \underline{\tau} = \tau \dot{i} = .002\rho_{air}(v_{wind} - u(0))\dot{i}$   $v_z = 0$  at z = 0(u, v) = 0 at  $z = \infty$ 

The Ekman layer thickness is constant in time and space since vortex diffusion balances the Coriolis force.

## Averages: For turbulent flow $\underline{V}(\underline{x}, t)$ , $p(\underline{x}, t)$ are random runctions of time and must be evaluated statistically using averaging techniques: time, ensemble, phase, or conditional. Time Averaging For stationary flow, the mean is not a function of time and we can use time averaging. (12) (18) $\overline{u} = \frac{1}{T} \int_{t_0}^{t_0+t} (t) dt \quad T > \text{any significant period of } u' = u - \overline{u}$ Fig. 6.1 Schematic representation of the triple decomposition: \_\_\_\_\_, instantaneous signal; ---, conditionally averaged; \_\_\_\_\_, mean. The numerals in circles characterize the phase (e.g. 1 sec. for wind tunnel and 20 min. for ocean) for later reference. Ensemble Averaging For non-stationary flow, the mean is a function of time and ensemble averaging is used $\overline{u}(t) = \frac{1}{N} \sum_{i=1}^{N} u^{i}(t)$ N is large enough that $\overline{u}$ independent $u^{i}(t) =$ collection of experiments performed under identical conditions (also can be phase aligned for same t=o). Statistically stationary process = the statistical properties, such as mean, variance and autocorrelation, do not change over time. Phase and Conditional Averaging Similar to ensemble averaging, but for flows with dominant frequency content or other condition, which is used to align time series for some phase/condition. In this (b) No (a) Stationary case triple velocity decomposition is used: u = u + u'' + u'Fig. 12.2 Stationary and where u'' is called organized oscillation. Phase/conditional averaging extracts all three components. Averaging Rules: $g = \overline{g} + g'$ $f = \overline{f} + f'$ s = x or t $\overline{\overline{f}} = \overline{f}$ $\overline{fg} = \overline{fg}$ $\overline{f'g} = 0$ $\overline{f}' = 0$ $\frac{\partial f}{\partial x} = \frac{\partial \overline{f}}{\partial x}$ $\overline{f+g} = \overline{f} + \overline{g}$ $\overline{fg} = \overline{fg} + \overline{f'g'}$ $\overline{\int f \, ds} = \int \overline{f} \, ds$ 10 AM 8 AM 9 AM Fig. 12.3 An ensemble of functions u(t).

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Statistical analysis needed for U, D= VXV,  
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Registles Decomposition  

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Homogeneous turbulence: the time-averaged properties of the flow are uniform and independent of position. For example, whereas  $\overline{u^2}$ ,  $\overline{v^2}$ ,  $\overline{w^2}$  may differ from each other, each must be constant throughout the system.

Isotropic turbulence: Turbulence in which the products and squares of the velocity components and their derivatives are independent of direction, or, more precisely, invariant with respect to rotation and reflection of the coordinate axes in a coordinate system moving with the mean motion of the fluid. Then all the normal stresses are equal, and the tangential stresses are zero.  $\overline{u^2} = i \overline{v^2} = i \overline{w^2}, \quad \overline{uv} = \overline{uw} = \overline{vw} = 0$ 

Isotropy implies homogeneity, but not vice versa.

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$$\lim_{x \to 0} z = \overline{z}(x) = \frac{z}{z} \xrightarrow{z} \qquad xssum_{x} = \overline{z}_{x}' = \overline{z}_{x}''$$

$$= \overline{z} \quad for \quad homogeneous \quad furtheme \\ urbane \quad \overline{z} = \underline{z} \quad uscledent$$

$$F_{y} = 22 \text{ Definition of the microscale}$$

$$\int_{0}^{1} \int_{0}^{1/2} \int_{0}^{0$$

More generally for two rembors revealed  

$$\pi_{i}(\underline{u}, \underline{e}_{i}) \perp \pi_{i}(\underline{u}, \underline{e}_{i}) = \pi_{i}(\underline{u}, \underline{e}_{i})\pi_{i}(\underline{v}, \underline{e}_{i})$$

$$R_{i}(\underline{u}, \underline{e}_{i}) \perp \pi_{i}(\underline{u}, \underline{e}_{i}) = \pi_{i}(\underline{u}, \underline{e}_{i})\pi_{i}(\underline{v}, \underline{e}_{i})$$

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$$R_{i}(\underline{u}, \underline{e}_{i}) \perp \pi_{i}(\underline{u}, \underline{e}_{i}) \perp \pi_{i}(\underline{v}, \underline{e}_{i})$$

$$R_{i}(\underline{u}, \underline{e}_{i}) = \pi_{i} > 0 \perp \pi_{i} = q_{i} \otimes 1 \quad \pi_{i}(\underline{u}, \underline{e}_{i}) \perp \pi_{i}(\underline{v}, \underline{e}_{i})$$

$$R_{i}(\underline{v}, \underline{e}_{i}) \geq 0 \quad \pi_{i} \geq 0 \quad \pi_{i} = q_{i} \otimes 1 \quad \pi_{i}(\underline{v}, \underline{e}_{i}) + \pi_{i} = \pi_{i} \otimes 1 \quad \pi_{i} \geq 0 \quad \text{second}$$

$$R_{i}(\underline{v}, \underline{v}) \perp \pi_{i} \geq 0 \quad \pi_{i} \geq 0 \quad \pi_{i} \geq 0 \quad \text{second}$$

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$$R_{i}(\underline{v}, \underline{e}_{i}) \geq 0 \quad \pi_{i} = \pi_{i} \times 1 \quad \pi_{i$$

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Spatial Spectra  
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$$\frac{e_{ij}}{e_{ij}} = \frac{1}{2} \sum_{i=1}^{i} e_{ij}} \frac{f(x)}{e_{ij}} \frac{e_{ij}}{e_{ij}} \frac{e_{ij}}{e_{ij}}$$

Calculate the surface area of a sphere with a radius of R:



Alternately, the area element can be written as  $dS(k) = k^2 \sin\varphi d\theta d\varphi$  according to the following figure.



Fine Speaks  
Fourier 
$$\hat{R}_{\mathcal{C}}(\omega') = \begin{bmatrix} e^{-i\varepsilon \omega'} \hat{R}_{\mathcal{C}}(z) \ dz \end{bmatrix} \begin{bmatrix} \mathbf{F}_{\mathcal{C}}(z) = \overline{u}(z) \ d(z+\tau) \end{bmatrix}$$
  
Transform  
Transform  
Prive  $\hat{R}_{\mathcal{C}}(z) = \frac{1}{2\pi} \int \hat{e}^{-i\varepsilon \omega'} \hat{R}_{\mathcal{C}}(\omega') \ d\omega' \qquad \omega' = 2\pi \omega \quad \omega \\ \hat{R}_{\mathcal{C}}(z) = \frac{1}{2\pi} \int \hat{R}_{\mathcal{C}}(\omega') \ d\omega' = \overline{z} = \int \hat{R}_{\mathcal{C}}(z\pi \omega) \ d\omega \\ \hat{R}_{\mathcal{C}}(z) = \frac{1}{2\pi} \int \hat{R}_{\mathcal{C}}(\omega') \ d\omega' = \overline{z} = \hat{R}_{\mathcal{C}}(z\pi \omega) \ d\omega \\ \hat{R}_{\mathcal{C}}(z) = \frac{1}{2\pi} \int \hat{R}_{\mathcal{C}}(\omega) \ d\omega = 1 \qquad = \hat{R}_{\mathcal{C}}(z\pi \omega) \ d\omega \\ \hat{R}_{\mathcal{C}}(z) = \frac{1}{2\pi} \int \hat{R}_{\mathcal{C}}(\omega) \ d\omega = 1 \qquad = \hat{R}_{\mathcal{C}}(z\pi \omega) \ d\omega \\ \hat{R}_{\mathcal{C}}(z) = \frac{1}{2\pi} \int \hat{R}_{\mathcal{C}}(\omega) \ d\omega = 1 \qquad = \hat{R}_{\mathcal{C}}(z\pi \omega) \ d\omega \\ \hat{R}_{\mathcal{C}}(z) = \frac{1}{2\pi} \int \hat{R}_{\mathcal{C}}(\omega') \ d\omega = 1 \qquad = \hat{R}_{\mathcal{C}}(z\pi \omega) \ d\omega \\ \hat{R}_{\mathcal{C}}(z) = \frac{1}{2\pi} \int \hat{R}_{\mathcal{C}}(\omega) \ d\omega \\ \hat{R}_{\mathcal{C}}(z) = \frac{1}{2\pi} \int \hat{R}_{\mathcal{C}}(\omega') \ d\omega \\ \hat{R}_{\mathcal{C}}(z) = \frac{1}{2\pi} \int \hat{R}_{\mathcal{C}}(\omega') \ d\omega' \\ \hat{R}_{\mathcal{C}}(z) = \hat{R}_{\mathcal{C}}(z) \ d\omega' \\ \hat{R}_{\mathcal{C}}(z) = \hat{R}_{\mathcal{C}}(\omega') \ d\omega' \\ \hat{R}_{\mathcal{C}}(\omega') \ d\omega' \\ \hat{R}_{\mathcal{C}}(z) = \hat{R}_{\mathcal{C}}(z) \ d\omega' \\ \hat{R}_{$